

**ABRASIVE NOTCHES ALONG THE ATLANTIC PATAGONIAN
COAST AND THEIR POTENTIAL USE AS SEA LEVEL MARKERS:
THE CASE OF PUERTO DESEADO (SANTA CRUZ, ARGENTINA)**

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3 1 ABRASIVE NOTCHES ALONG THE ATLANTIC PATAGONIAN COAST AND THEIR
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5 2 POTENTIAL USE AS SEA LEVEL MARKERS: THE CASE OF PUERTO DESEADO
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7 3 (SANTA CRUZ, ARGENTINA)
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26 11 ABSTRACT

27 12 Patagonia Argentina is a key area for the study of sea level changes in the Southern
28 13 hemisphere, but the availability of reliable sea level markers in this area is still problematic.

29 14 In fact the storm deposits (beach ridge) commonly used here to reconstruct past sea level
30 15 oscillations introduce a wide error. Along the Puerto Deseado coast (Santa Cruz),
31 16 morphometric analyses of 11 features were carried out using traditional measurement
32 17 tools and a digital software-based method (tested on one selected feature) with the aim to
33 18 investigate the possibility of their use as sea-level markers. By undertaking accurate
34 19 topographic profiles we identified the relationship between notches and current sea level.

35 20 In detail, we identified two clusters of notch retreat point elevations, with a very low internal
36 21 variability. The lower was located a little below the mean high-tide level (mHT) and the
37 22 upper located at least about 0.5 m above the Maximum high-tide level (MHT). Field
38 23 observations of tidal levels and notches position suggest that the lower notches are active
39 24 and the upper are inactive. This study on the abrasive notches attests their quality as sea
40 25 level markers and opens up the use of fossil abrasive notches as palaeo sea-level markers
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3 26 because the error linked to these features is substantially smaller than that introduced by
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5 27 beach ridges commonly used in the study area.
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16 29 Key words: notch, rocky coast, sea level marker, Patagonia, Argentina.
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29 31 INTRODUCTION

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32 Patagonia offers a unique opportunity in the South Hemisphere to study the
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34 palaeoenvironmental evolution of an area that lies between the mid and the low latitudes
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36 from the tropics to the margins of Antarctica. As a passive margin this area is tectonically
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38 relatively stable and is particularly suitable for the study of the eustatic sea-level
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40 oscillations over the Late Pleistocene and Holocene (Milne et al. 2005; Milne and Mitrovica
41
42 2008).The Patagonian coast is mainly characterized by a meso-macrotidal regime (Isla
43
44 and Bujalesky, 2008) that in itself makes the study of palaeo sea-level difficult. Moreover,
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46 in Patagonia most of the studies on this topic are based on coastal deposits such as beach
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48 ridges (Codignotto et al., 1992; Rostami et al., 2000; Schellmann and Radtke, 2000; 2003;
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50 2010; Pedoja et al., 2011; Ribolini et al., 2011; Isola et al., 2012; Zanchetta et al., 2012;
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52 Pappalardo et al. 2014). These deposits depending on the storm energy, wave exposure
53
54 and tide range, usually lead to an overestimation of the mean sea-level value (e.g.
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56 Schellmann and Radtke, 2010). The best available deposits used to reconstruct palaeo
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58 sea-level in the Patagonian coast are river mouth terraces, formed by the interfingering of
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60 fluvial and littoral sediments at the outlet of rivers into the ocean. Their elevation, in fact,
could be directly related to the high tide level (Schellmann and Radke, 2010), but their
contribution to the study of past sea level is limited because they are very limited in
occurrence. Coastal deposits and their significance as sea level markers have been
largely studied along the Patagonian coast, while rocky coast have been nearly ignored, in
spite of the presence of rock outcrops in coastal areas and the occurrence of notches.

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3 52 These markers have been used in Puerto Deseado area by Pedoja et al. (2011) to state
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5 53 Pleistocene sea levels; they assumed for these features a direct connection with mean sea
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7 54 level. When used in order to assess Holocene palaeo sea levels in meso-macrotidal areas,
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9 55 though, the degree of accuracy required implies that their occurrence at a specific
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11 56 elevation within the tidal-range should be stated exactly.

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14 57 In addition, abrasive notches in meso and macrotidal areas are poorly studied around the
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16 58 world probably because the tidal regime and the genetic typology of notches make difficult
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18 59 to understand their significance as sea level markers. In fact, notch shape and its
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20 60 relationship with sea-level is strictly linked to the predominant genetic process. Tidal
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22 61 notches formed in carbonatic rocks, by bioerosion and carbonate dissolution, are widely
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24 62 used and considered one of the best mean sea level indicators (i.e. Pirazzoli, 1986; Rust
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26 63 and Kershaw 2000; Kershaw and Guo, 2001; Lambeck et al. 2004; Kelletat, 2005;
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28 64 Evelpidou et al., 2012; Moses, 2013). Even if a debate on their genesis is still open
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30 65 (Moses, 2013), in microtidal area such as the Mediterranean sea, the elevation of the
31
32 66 retreat point (i.e. the apex or the maximum curvature point of the notch profile) is
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34 67 considered to correspond to the palaeo mean sea-level with an accuracy that in the best
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36 68 cases reaches ± 0.1 m (Ferranti et al., 2006). On the contrary, notches developed in non
37
38 69 carbonatic rock by a predominantly abrasive process are generally considered unreliable
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40 70 as sea level marker. Commonly the relation between this type of notch and sea level is
41
42 71 considered changing according to the availability of abrasive sediments and wave energy
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44 72 (Pirazzoli, 1986, Kelletat 2005). Nevertheless a morphometric study by Trenhaile et al.
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46 73 (1998) in the Canadian Coast attests that the retreat point of abrasive notches in macro
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48 74 tidal area is located a little below the mean high tide level. Dickinson (2000) in the Pacific
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50 75 Islands observed that the retreat point is located close to the high tide level, while Irion et
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52 76 al. (2012) related the floor of the notches to the spring high tide in the Brazilian Coast. This
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54 77 work describes abrasive notches, carved in volcanic massive rocks and their relation with
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3 78 current sea level in the meso-macrotidal environment of Puerto Deseado area. The study
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5 79 explores also the possibility to extrapolate this correlation to relict notches, and assesses
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7 80 their utility for measuring past sea levels. To understand the relationship between sea
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9 81 level and abrasive notches would improve their use as sea level marker, and contribute to
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11 82 the reconstruction of sea-level variations in the recent past in a key area for the study of
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13 83 climate change.
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17 85 THE STUDY AREA

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23 87 During the field surveys in 2010 and 2011, notches were found in different parts of the San
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25 88 Jorge Gulf (44° - 47° S Santa Cruz-Chubut Provinces), particularly in the Bahía
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27 89 Bustamante, Camarones and Deseado areas. Due to the availability of astronomical data
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29 90 we chose to concentrate our work in the southern portion of the gulf, in the Puerto
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31 91 Deseado area (Fig. 1).
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34 92 Local climate is characterized by a mean annual precipitation around 200 mm and an
35
36 93 average annual air temperature of 8.2°C (Servicio Meteorológico Nacional-Argentina,
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38 94 <http://www.smn.gov.ar>). The coastal area experiences a meso-macrotidal regime (>4
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40 95 mlsla and Bujalesky, 2008), with a tide range usually fluctuating between 2.5 and 5.5
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42 96 metres (http://www.hidro.gov.ar/Oceanografia/Tmareas/Form_Tmareas.asp). The lunar
43
44 97 and solar cycles produce a mixed tidal type yielding unequal low and high tides (Fig. 2).
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47 98 The analysed notches are located along a W-E oriented coastline (Fig. 1), impacted by
48
49 99 prevailing winds blowing from the West, less frequent are winds blowing from the
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51 100 Southwest and the North (www.windfinder.com).
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54 101 The study area is located in the Deseado massif geological province, characterized by the
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56 102 outcrop of middle-upper Jurassic volcanic and volcanoclastic rocks of the Chon-Aike
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58 103 Formation of Bahía Laura Group (Guido et al., 2004; Sruoga et al. 2004). These rocks are
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3 104 mainly calc-alkaline and peraluminous rhyolites, with some dacitic members (Pankhurst
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5 105 and Rapela, 1995; Pankhurst et al., 1998), displaying a prevalent massive aspect
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7 106 (pyroclastic density currents deposits). The principal minerals are quartz, K-feldspar,
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9 107 plagioclase and biotite; accessories are magnetite, ilmenite, apatite and zircon with
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11 108 associated monazite (Sruoga et al. 2004).

13
14 109 The volcanic and volcanoclastic substratum dominates the area of Puerto Deseado
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16 110 settlement and part of the Rio Deseado river mouth. It is shaped in the form of rocky
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18 111 marine terraces sparsely covered by Quaternary marine and subaerial deposits (Ribolini et
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20 112 al. 2014; Zanchetta et al. 2014). A rocky promontory (Punta Foca-Punta Cavedish), with
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22 113 an indentation of a gravelly pocket beach, separates the Puerto Deseado area from the
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24 114 northern coast (Fig. 1). On the northern coast there are littoral deposits of Quaternary age
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26 115 that seaward cover a modern intertidal-lower supratidal shore platform (Feruglio, 1950;
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28 116 Codignotto et al., 1988; Zanchetta et al. 2014).

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33 34 118 METHOD

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38 120 Notches present in this area show a wide variability of dimensions and shapes (Fig. 3).
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40 121 The studied notches are located parallel to the current coastline at different elevations
41
42 122 above HT level. The observed notches do not show evidence of structural control in their
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44 123 formation. In fact, they are carved in massive lithified pyroclastic units and there is no
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46 124 evidence of layering, discontinuity, fractures and structural weakness directions coinciding
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48 125 with the position of notches. The azimuthal distribution of the few detectable fractures in
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50 126 the surrounding area shows a maximum peak around N315E (Fig. 4), while the coastline
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52 127 directions shows a very narrow peak at N70E. Consequently no structural control in the
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54 128 development of notches could be supposed.
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3 129 The morphometry of these notches is described according to the parameters suggested by
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5 130 Pirazzoli (1986, Fig. 5) total height (H), distance of retreat point from cliff face (D), height of
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7 131 retreat point from the floor (H_f) and measured using a graduated bar.

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9 132 Taking into account that these parameters may vary inside a single feature depending on
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11 133 how they are measured a more accurate analyses was undertaken in one selected notch.

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13 134 This selected notch was surveyed using a software-based technology that allows the user

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15 135 to generate three-dimensional data from a surface using a series of overlapping images

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17 136 taken by a consumer-grade digital camera. In this case, 30 images have been processed

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19 137 by the open-source SfM software package Bundler (Snavely et al., 2006; 2007), obtaining

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21 138 a sparse 3D coloured point cloud with internally-consistent 3D geometry (Fig. 6). A

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23 139 detailed description of this method can be found in Favalli et al., (2011). The 3D point

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25 140 cloud has been processed using the free MeshLab software

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27 141 (<http://meshlab.sourceforge.net/>), which allows the editing process of the points. The post-

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29 142 processing phase was done in a GIS environment selecting the points belonging to the

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31 143 water surface filling the pool in front of the notch, then rotating the whole point cloud so

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33 144 that the selected pool points lay on a horizontal plane. Multiview 3D reconstructions have

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35 145 root mean square (RMS) errors that are usually less than 1% of the average linear

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37 146 dimensions. In our case the RMS error is under 5 cm (e.g. Cecchi et al., 2003; Favalli et

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39 147 al., 2011). After that 13 profiles, spaced at 30 cm intervals, were derived (Fig.7). The H, D,

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41 148 H_f parameters and the range of elevation of the retreat points above the horizontal plane

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43 149 have been calculated for each profile.

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45 150 The elevation above sea level was measured by means of graduated stacks equipped with

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47 151 a spirit level (three-meter rod). Through a centimetre-graduated bar, we constructed a

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49 152 topographic profile between the sea level and the retreat point of each notch, and we

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51 153 corrected the elevation data accordingly to the tide level at the time and date of

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53 154 measurement derived from astronomic tide tables, calculated for the 2011/01/28 in Puerto

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3 155 Deseado. Through this method we reconstructed the topography of the shore platform in
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5 156 front of the notches, the elevation of morphological markers of high tide and the elevation
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7 157 of cliff at centimetric resolution.
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10 158 The appearance of surface (polished, smooth etc.) and the lateral continuity of the notches
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12 159 were qualitatively described to better understand their genesis. We also analysed the
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14 160 relationship of the notches with other geomorphological and biological sea-level markers
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16 161 occurring in the same area as the high-tide morphological step in the beach, the upper
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18 162 limit of the ecological zone of barnacles, mussels and green-seaweed.
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22 164 RESULTS

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27 166 The shapes and morphometric characteristics of 11 notches directly measured in the field
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29 167 are reported in Figure 8. In particular, the height (H) ranges between 90 cm and 260 cm,
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31 168 and the distance (D) ranges between 21 and 1020 cm. All studied notches display an
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33 169 asymmetric shape with a gently sloping floor rising up to the retreat point that is quite low
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35 170 as shown by the small values of the elevation difference between retreat point and the
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37 171 floor (H_f).
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40 172 The retreat point area is characterized by an abrupt decrease in roughness, with a rock
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42 173 surface notably smoother on touch than the notch roof and floor (Fig.3). Most notches
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44 174 show potholes located in the inner part of the floor, just below the retreat point, where
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46 175 abrasive materials were trapped against the foot of scarps. These potholes are
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48 176 characterized by elliptic shape with main axis coincident with notch length ranging from 0.4
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50 177 m to 5 m, minor axis ranging from 0.1 to 0.6 m, and about 0.3 m deep. (Fig.3). Potholes
51
52 178 are generally colonized by mussels (*Brachidontes purpuratus*), barnacles (*Balanus* sp.),
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54 179 and green algae. Marine water usually remains inside the potholes even during low tide
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56 180 (Fig.3).
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3 181 Contrasting colours belts evidence the vertical biological zonation on the rock, in
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5 182 accordance with Stephenson and Stephenson (1972). This zonation is typical of sheltered
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7 183 areas (Stiros and Pirazzoli, 2008) and results from changes in the biological community
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9 184 structure. Specifically, from base to top of the zonation, the upper boundary of green
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11 185 algae marks the mean low tide at very low tide conditions, followed by a black band of
12
13 186 mussels, white barnacles and black lichens. The continuous barnacle band upper limits
14
15 187 marks the minimum high tide level in sheltered zone (Little et al. 2009; Stiros and Pirazzoli,
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17 188 2008). The retreat points of notches 3, 4, 5, 7 and 8 (Fig.8) are located just above the band
18
19 189 characterized by mussels and barnacles where abrasion prevents colonization of rock
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21 190 surface by organisms. Lichens are visible up to 1 m above the retreat point. On the
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23 191 contrary, the other notches do not show biological colonization.
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27 192 A rocky shore platform is present in front of most of the notches, patchily covered by mixed
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29 193 sediment including sand and eotometric pebbles (from centimetre to decimetre in
30
31 194 maximum diameter) of different lithologies. They are repeatedly moved by wave action as
32
33 195 to excavate a great number of smoothly circular potholes on the rocky platform, attesting
34
35 196 the potency of wave abrasion (Trenhaile 2004). The basal part of notches 6, 10 and 11
36
37 197 (Fig.8) is partially buried by coarse deposits. The deposit burying notch 6 consists on top
38
39 198 of a few centimetres of a pebble lag with no evidence of soil development, covering ca 10
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41 199 cm of dark, charcoal- and ash-rich sandy deposit containing abundant shell material and
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43 200 flints. This deposit is certainly an archaeological layer, covering the undisturbed beach
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45 201 ridge deposits below.
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49 202 To better constrain the parameters variability within the same feature, a 3D digital model of
50
51 203 notch 1 was derived using the multiview reconstruction (Favalli et al., 2011). Then 13
52
53 204 strips were generated (Fig. 7b) and the morphometric parameters calculated for each strip
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55 205 are listed in Table 1. Most of the calculated values have a large variability, in fact, H varies
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57 206 between about 180 and 300 cm, while D varies between about 80 and 320 cm. The
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3 207 platform angle in front of the notch varies between about 3° to 7° (the higher values are
4
5 208 related to the seaward slope of the pothole located in the notch floor). The variability of the
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7 209 retreat point elevations referenced to a horizontal plane is instead very low, ca. 30 cm.

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9 210 Six topographic profiles (Fig.9) were undertaken in order to investigate the relationship
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11 211 between notch elevation, sea level morphological indicators (fair weather berm, seaweed
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13 212 beach wrack from the previous high tide) and high-tide elevation derived from astronomic
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15 213 tide tables. All profiles show a good correlation between the elevation of the observed high
16
17 214 tide evidence and the corresponding value derived from tide table. The retreat points
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19 215 elevation ranges between -30 cm and +170 cm with respect to the high tide. Based on
20
21 216 these elevations, notches could be split in two different groups: the first including notches
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23 217 3, 4, 5, 7, 8 and 10 located close to high tide (HT) level, the second including notches 1, 2
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25 218 ,6, 9 and 11 located about 1.5 m above HT. Direct observations during the days of the field
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27 219 work confirm that water rises up to the first group during HT.
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221 DISCUSSION

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223 There is a large variability in the morphometric parameters calculated among the studied
224 notches (Fig.8) and also in those referred to a single feature (e.g. notch 1) (Fig. 7a, b).
225 Nevertheless, morphometric analysis of notch 1 shows that the retreat point elevation has
226 a very low variability (Fig.7b). This make us confident in using the value of retract point
227 elevation for sea level estimation.

228 The topographic profiles carried out to calculate the elevation of retreat points with respect
229 to HT of the 2011-01-28 provided two groups of elevation values, the first approximately
230 coincident with the high tide level, the second ranging about 1.50 m above it (Fig. 9). The
231 variability of retreat point elevation shows two clusters with a very low internal variability (\pm
232 0.3 m for the lower and \pm 0.2 m for the upper notches, Fig. 10). When compared with the

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3 233 tide curves derived from the local tide tables (Fig. 11), the elevation of lower group of
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5 234 notches is consistent with the monthly mean value of high tide (mHT), calculated on the
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7 235 tide table values recorded during the 15 months considered (Fig. 11). The elevation of
8
9 236 upper notches is located at least ~0.5 m above the curve of the highest astronomical tide
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11 237 (MHT), consequently they can be reached by waves only during storms. Furthermore, at
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13 238 the base of notch 6 (upper notches group) we found a shell midden deposit radiocarbon
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15 239 dated at 0-236 cal. yr B.P. (Table 2) (*Mytilus edulis*), and at the base of the sediment
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17 240 covering notch 11 (upper notches group) we dated an individual of *Nacella deaurata*
18
19 241 yielding a radiocarbon age 476-521 cal yr B.P. (Zanchetta et al., 2014). The preservation
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21 242 of these deposits suggests that, even if occasionally reached by waves, the rates of
22
23 243 erosion are so low that it cannot remove an unconsolidated deposit. These observations,
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25 244 along with the tidal values, allow us to consider the lower notches group as active forms
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27 245 and the upper ones as inactive. The presence/absence of organisms related to sea level in
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29 246 the lower/higher notches group supports this hypothesis. The limited elevation range of
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31 247 retreat point inside each group enable us to consider abrasive notches as a good
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33 248 morphological sea-level marker in this bay. Moreover the retreat point elevation of current
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35 249 notches in comparison with the high-tide level variation suggests that they mark the
36
37 250 medium level of high tide. These results support the suitability of abrasive notches as
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39 251 palaeo sea-level markers in this area.

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41 252 Given that their elevation is linked to the mean high tide level, they display a small error
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43 253 (± 0.30 m) when compared with other sea level markers commonly used in the Patagonian
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45 254 coast, as beach ridge crests, that overestimate the mean sea level value. Moreover,
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47 255 because of their genesis, notches indicate a marine stillstand phase while the other marine
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49 256 deposits commonly used in this area form during regressive phases (Pirazzoli, 1986; 1996;
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51 257 Kershaw and Guo, 2001; Kelletat, 2005; Ramos and Tsutsumi, 2010). On the other hand,
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53 258 marine deposits are suitable for dating because of their fossil content, while notches are
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3 259 quite difficult to date. Therefore a combined use of marine deposit chronology and related
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5 260 elevation of coeval notch may offer a more confident reconstruction of the past sea level.
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7 261 The study area is a pocket beach characterized by the presence of several well-defined
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9 262 beach ridge crests testifying different phases of coastal aggradation (Zanchetta et al.,
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11 263 2014). Because of their position and the associated deposits, the notches with elevation
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13 264 ca. 1.5 m above current mean high tide can be correlated with the lowermost inactive
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15 265 beach ridges visible on the shoreplain located at ca 3 m above high tide (Zanchetta et al.,
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17 266 2014). This indicates that in a such protected bay the height of the beach ridge crest
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19 267 exceeds of ca 1.5 m that of notches i.e. the mean High Tide level. This datum is confirmed
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21 268 by our field measurements on active beach ridge crests attesting their elevation at about
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23 269 1.3 – 1.5 above high tide.
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271 CONCLUSION

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273 Along the meso-macro tidal coast of Patagonia (Puerto Deseado area), the retreat point of
274 active abrasive notches approximates the medium level of high tide with a small variability
275 (± 0.3 m). The retreat point seems not only correlated to the high-tide level, as already
276 observed by Dickinson (2001) in the Pacific Islands, and Irion et al. (2012) on the coast of
277 Brazil, but more precisely it is located slightly below the medium value of high tide level, in
278 perfect agreement with the study made by Trenhaile et al. (1998) on the New Brunswick
279 coast (Canada). This coherence between our data and those obtained in different
280 environmental conditions underlines the general significance of the results obtained in this
281 part of the Patagonian coast, excluding the influence of local factors. This result opens to
282 the possibility of using fossil abrasive notches as palaeo sea-level markers, more effective
283 than the markers frequently adopted in this coastal region (i.e. beach ridge crest,
284 mouth/littoral terrace). Indeed, the lowermost inactive notches in this area is at $1.5 \text{ m} \pm 0.2$

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3 285 m above medium level of high tide, whereas the coeval beach ridge crest stands at 3 m
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5 286 above high tide.

7 287 An accurate determination of relative sea level through the selection of the appropriate
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9 288 markers such as notches is particularly useful to test geophysical models and to
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11 289 disentangle the glacio-eustatic, hydro-isostatic and neotectonic components in sea level
12
13 290 change signal, a topic still debated in the Patagonian Coast.

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311 References

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313 Cecchi E, van Wyk de Vries B, Lavest JM, Harris A, Davies M. 2003. N-view
314 reconstruction: a new method for morphological modelling and deformation
315 measurement in volcanology. *Journal of Volcanology and Geothermal Research*,
316 **123**: 181–201.

317 Codignotto JO, Marcomini SC, Santillana SN. 1988. Terrazas marinas entre Puerto
318 Deseado y Bahía Bustamante, Santa Cruz, Chubut. *Asociación Geológica*
319 *Argentina* **43**: 43–50.

320 Codignotto JO, Kokot RR, Marcomini SC. 1992. Neotectonism and sea level changes in
321 the coastal zone of Argentina. *Journal of Coastal Research* **8**: 125–133.

322 Dickinson WR. 2000. Hydro-isostatic and tectonic influences on emergent Holocene
323 paleoshorelines in the Mariana Islands western Pacific Ocean. *Journal of Coastal*
324 *Research*, **16**: 725–746

325 Dickinson WR. 2001. Paleoshoreline record of relative Holocene sea levels on Pacific
326 islands. *Earth-Science reviews* **55**: 191–234.

327 Evelpidou N, Kampolis I, Pirazzoli PA, Vassilopoulos A. 2012. Global sea-level rise and
328 the disappearance of tidal notches, *Global and Planetary Change*, **92–93**: 248–256

329 Feruglio E. 1950. Descripción geológica de la Patagonia. *Dirección General de Y.P.F.* **3**:
330 74-197, Buenos Aires.

331 Favalli M, Fornaciai A, Isola I, Tarquini S, Nannipieri M. 2011. Multiview 3D reconstruction
332 in geosciences. *Computer & Geosciences*, **44**: 168–176.

333 Ferranti L, Antonioli F, Mauz B, Amorosi A, Dai Pra G, Mastronuzzi G, Monaco C, Orru P,
334 Pappalardo M, Radtke U, Renda P, Romano P, Sansò P, Verrubbi V. 2006.
335 Markers of the last interglacial sea-level high stand along the coast of Italy: Tectonic
336 implications. *Quaternary International* **145–146**: 30–54.

- 1
2
3 337 Guido D, Escayola M, Schalamuk I. 2004. The basement of the Deseado Massif at Bahía
4
5 338 Laura, Patagonia, Argentina: A proposal for its evolution. *Journal of South -*
6
7 339 *American Earth Science*, **16** (7): 567–577.
- 8
9
10 340 Hughen KA, Baillie MGL, Bard E, Bayliss A, Beck JW, Bertrand C, Blackwell PG, Buck CE,
11
12 341 Burr G, Cutler KB, Damon PE, EdwardsRL, FairbanksRG, Friedrich M, Guilderson
13
14 342 TP, Kromer B, McCormac FG, Manning S, Bronk Ramsey C, Reimer PJ, Reimer
15
16 343 RW, Remmele S, Southon, JR, Stuiver M, Talamo S, TaylorFW, Van der Plicht J.,
17
18 344 WeyhenmeyerCE. 2009. Marine04 Marine radiocarbon age calibration, 26 - 0 ka BP.
19
20 345 *Radiocarbon* **46**: 1059–1086
- 21
22
23 346 Irion G, Onofre de Morais J, Bungenstock F. 2012. Holocene and Pleistocene sea-level
24
25 347 indicators at the coast of Jericoacoara, Ceará, NE Brazil, *Quaternary Research*, **77**
26
27 348 (2): 251–257.
- 28
29
30 349 Isla FI, Bujalesky GG. 2008. Coastal Geology and Geomorphology of Patagonia and
31
32 350 Tierra del Fuego Archipelago. In *The late Cenozoic of Patagonia and Tierra del*
33
34 351 *Fuego*. Rabassa J (ed.) Elsevier Science Publishers: Amsterdam; 227-240.
- 35
36 352 Isola I., Bini M., Ribolini A., Pappalardo M., Consoloni I., Fucks E., Boretto G., Ragaini L.,
37
38 353 Zanchetta G. 2011. Geomorphologic map of northeastern sector of san jorge gulf
39
40 354 (chubut, argentina) *Journal of Maps*, 476-485.
- 41
42
43 355 Kelletat DH. 2005. Notches. In *Encyclopedia of coastal science*. Edited by Schwartz
44
45 356 Springer.
- 46
47 357 Kershaw S, Guo L. 2001 Marine notches in coastal cliffs: indicators of relative sea level,
48
49 358 Peracora Peninsula, central Greece. *Marine Geology* **179** (3-4): 213–228.
- 50
51
52 359 Lambeck K, Antonioli F, Purcell A, Silenzi S. 2004. Sea level change along the Italian
53
54 360 coast for the past 10,000 yrs. *Quaternary Science Reviews* **23**:1567-1598.
- 55
56 361 Little C, Williams GA, Trowbridge CD. 2009. The biology of rocky shores. Oxford
57
58 362 University Press: Oxford
- 59
60

- 1
2
3 363 Milne GA, Long AJ, Bassett SE. 2005. Modelling Holocene relative sea-level observations
4
5 364 from the Caribbean and South America. *Quaternary Science Reviews* **24**: 1183–
6
7 365 1202.
8
9 366 Milne GA, Mitrovica JX. 2008. Searching for eustasy in deglacial sea-level histories.
10
11 367 *Quaternary Science Reviews* **27**: 2292–2302.
12
13 368 Moses, C. A. 2013. Tropical rock coasts Cliff, notch and platform erosion dynamics.
14
15 369 *Progress in Physical Geography*, **37**(2): 206-226.
16
17 370 Pappalardo M, Aguirre M, Bini M, Consoloni I, Fucks E, Hellstrom J, Isola I, Ribolini A
18
19 371 Zanchetta G. 2014. Coastal landscape evolution and sea-level change:a case study
20
21 372 from Central Patagonia (Argentina). *Zeitschrift für Geomorphologie*, DOI:
22
23 373 10.1127/0372-8854/2014/0142.
24
25 374 Pankhurst R, Rapela C. 1995. Production of Jurassic rhyolite by anatexis of the lower crust
26
27 375 of Patagonia. *Earth and planetary Science Letters* **134**: 23–36.
28
29 376 Pankhurst R, Leat P, Sruoga, P, Rapela C, Márquez M, Storey B, Riley T. 1998. The Chon
30
31 377 Aike province of Patagonia and related rocks in West Antarctica: a silicic large
32
33 378 igneous province. *Journal of Volcanology and Geothermal Research* **81**: 113–136.
34
35 379 Pedoja K, Husson L, Regard V, Cobbold PR, Ostanciaux E, Johnson ME, Kershaw S,
36
37 380 Saillard M, Martinod J, Furgerot L, Weill P, Delcaillau B. 2011. Relative sea-level fall
38
39 381 since the last interglacial stage: Are coasts uplifting worldwide? *Earth-Science*
40
41 382 *Reviews* **108** (1-2): 1–15.
42
43 383 Pirazzoli PA. 1986. Marine notches. In *Sea level research: a manual for the collection and*
44
45 384 *evaluation of data*. O. Van de Plassche (ed) Geo Books: Norwich
46
47 385 Pirazzoli PA. 1996. *Sea-Level Changes: The Last 20,000 Years*, Wiley. John Wiley &
48
49 386 Sons: Chichester.
50
51
52
53
54
55
56
57
58
59
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- 1
2
3 387 Ramos NT, Tsutsumi H. 2010. Evidence of large prehistoric offshore earthquakes deduced
4
5 388 from uplifted Holocene marine terraces in Pangasinan Province, Luzon Island,
6
7 389 Philippines. *Tectonophysics* **495**: 145–158.
- 9
10 390 Ribolini A, Aguirre M, Baneschi I, Consoloni I, Fucks E, Isola I, Mazzarini F, Pappalardo M,
11
12 391 Zanchetta G, Bini M. 2011. Holocene beach ridges and coastal evolution in the
13
14 392 Cabo Raso bay (Atlantic Patagonian coast, Argentina). *Journal of Coastal*
15
16 393 *Research*, **27** (5): 973–983.
- 18
19 394 Ribolini A, Bini M, Consoloni I; Isola I, Pappalardo M, Zanchetta G, Fucks E, Panzeri L,
20
21 395 Martini M, Terrasi F. 2014. Late Pleistocene sand wedges along the patagonian
22
23 396 coast (Argentina): chronological constrains and implication for coastal permafrost
24
25 397 distribution. *Geografiska Annaler, Series A: Physical Geography* , DOI:
26
27 398 10.1111/geoa.12038
- 29
30 399 Rostami K, Peltier WR, Mangini A. 2000. Quaternary marine terraces, sea-level changes
31
32 400 and uplift history of Patagonia, Argentina: comparisons with predictions of the ICE-
33
34 401 4G (VM2) model of the global process of glacial isostatic adjustment. *Quaternary*
35
36 402 *Science Reviews* **19**:1495–1525.
- 38
39 403 Rust D, Kershaw S. 2000. Holocene tectonic uplift patterns in northeastern Sicily: evidence
40
41 404 from marine notches in coastal outcrops. *Marine Geology* **167** (1-2): 105–126.
- 42
43 405 Schellmann G, Radtke U. 2000. ESR dating stratigraphically well-constrained marine
44
45 406 terraces along the Patagonian Atlantic coast (Argentina). *Quaternary International*
46
47 407 **68/71**:261–273.
- 49
50 408 Schellmann G, Radke U. 2003. Coastal terraces and Holocene sea-level changes along
51
52 409 the Patagonian Atlantic coast. *Journal of Coastal Research* **19**: 963–996.
- 54
55 410 Schellmann G, Radtke U. 2010. Timing and magnitude of Holocene sea-level changes
56
57 411 along the middle and south Patagonian Atlantic coast derived from beach ridge
58
59
60

- 1
2
3 412 systems, littoral terraces and valley-mouth terraces. *Earth-Science Reviews Earth-*
4
5 413 *Science Reviews* **103**:1–30.
6
7 414 Snaveley N, Seitz SM, Zeliski RS. 2006. Photo tourism: exploring image collections in 3D.
8
9 415 SIGGRAPH conference proceedings. ACM Press: New York.
10
11 416 Snaveley N, Seitz D, Zeliski R. 2007. Modeling the world from internet photo collections.
12
13 417 *International Journal of Computer Vision* **80**: 189–210.
14
15
16 418 Sruoga, P., Rubinstein, N., Hinterwimmer, G. 2004. Porosity and permeability in volcanic
17
18 419 rocks: A case study on the Serie Tobifera, South Patagonia, Argentina. *Journal of*
19
20 420 *Volcanology and Geothermal Research*, **132** (1): 31–43.
21
22
23 421 Stephenson TA, Stephenson A. 1972. Life Between Tidemarks on Rocky Shore. WH
24
25 422 Freeman & Co, San Francisco.
26
27 423 Stiros SC., Pirazzoli PA. 2008. Direct determination of tidal levels for engineering
28
29 424 applications based on biological observations. *Coastal Engineering*, **55** (6): 459–
30
31 425 467.
32
33
34 426 Trenhaile AS, Pepper DA, Trenhaile RW, Dalimonte M. 1998. Stacks and notches at
35
36 427 Hopewell rocks, New Brunswick, Canada. *Earth Surface Processes and*
37
38 428 *Landforms. Earth Surf. Process. Landforms* **23**: 975–988.
39
40
41 429 Trenhaile, AS 2004. Modeling the accumulation and dynamics of beaches on shore
42
43 430 platforms. *Marine Geology* **206**: 55–72.
44
45 431 Zanchetta G, Consoloni I, Isola I, Pappalardo M, Ribolini A, Aguirre M, Fucks E, Baneschi
46
47 432 I, Bini, M, Mazzarini F, Ragaini L, Terrasi F. 2012. New insight on the Holocene
48
49 433 marine transgression in the Bahia Camarones (Chubut, Argentina). *Italian Journal*
50
51 434 *of Geosciences* **131** (1):19–31.
52
53
54 435 Zanchetta G., Bini M., Isola I., Pappalardo M., Ribolini A., Ragaini L. Terrasi F., Consoloni
55
56 436 I., Borretto G., Fucks E. 2014. Middle to Late Holocene relative sea level changes at
57
58 437 Puerto Deseado (Patagonia, Argentina). *Holocene*, **24** (3): 307–317.
59
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7 440 Fig. 1. The study area and the location of the 11 abrasive notches analysed (red points).
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11 442 Fig. 2. Tidal trends during the 15 months from 10/10 to 12/11
12 (http://www.hidro.gov.ar/Oceanografia/Tmareas/Form_Tmareas.asp). The lunar and solar
13 443 cycles produces in this area a mixed tidal regime yielding unequal low tide and high tide
14 444 (Max high tide; Max low tide; min high tide and min low tide). Elevation data are in
15 445 reference to the reduction plane (theoretical plane located under the mean sea level in
16 446 order to have only positive tidal values in the tables) located 3.20 m below the medium
17 447 level.
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450 Fig. 3 Examples of the observed notches
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452 Fig. 4 Analysis of azimuth distribution of fractures (a) and coastal line (b) directions
453 Notice how the predominant fracture direction is approximately at a right angle to the
454 coastline.
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456 Fig.5 Measured parameters: total height (H), distance of retreat point from cliff face (D),
457 height of retreat point from the floor (H_f), according to Pirazzoli (1986)
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459 Fig.6 a): Notch1 photography, b): 3D model displayed by MeshLab software, c): zoomed in
460 area on the point cloud.
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462 Fig.7 Variability of notches shape, shown through the 3D analysis modelling, a) 3D view of
463 the point cloud, b) the 13 derived profiles.
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5 465 Fig. 8. Morphometric characteristics of the 11 notches analyzed R_p = retreat point; H_f =
6 elevation of retreat point above the floor; H = notch height; D = notch depth; h_{TE} =
7 elevation of retreat point above high tide. All data are reported in cm and are derived from
8 field measurements.
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16 470 Fig. 9. Topographic profiles in which notch position (black arrow) is compared with
17 morphological evidence of high tide (black dashed arrow) and high tide elevation derived
18 from tide tables, calculated for the 2011/01/28 in Puerto Deseado (black dashed line).
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25 474 Fig. 10 Distribution of the retreat point elevations above high tide. In spite the low number
26 of features they show two well separated clusters with a very low variability.
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31 477 Fig. 11. Relation between retreat point elevation and tide level derived from the current
32 tide tables of Puerto Deseado. Elevation data are in reference to the reduction plane
33 (theoretical plane located under the mean sea level in order to have only positive tidal
34 values in the tables) located 3.20 m below the medium level.
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strip	Ht (cm)	H _f (cm)	D (cm)
2	266.4	35.7	88.6
4	230.7	21.4	77.1
6	239.3	22.1	92.1
8	238.6	32.9	137.1
10	247.1	12.1	172.9
12	272.9	26.4	202.9
14	211.4	22.1	192.9
16	303.6	43.6	322.9
18	284.3	13.6	301.4
20	260.0	45.0	255.0
22	237.1	52.9	217.9
24	225.7	25.7	181.4
26	185.7	20.7	138.6
Field	260	50	350

Sample	Field code	¹⁴ C yr BP (±1σ)	¹⁴ C cal yr BP (±1σ)	Species or material
DSH1959	WP344	880±30	476-521	<i>Nacella</i> <i>(Patinigera)</i> <i>deaurata</i>
DSH2735	WPi 39	480±60	0-236	<i>Aulacomya atr</i>

For Peer Review

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Fig. 1. The study area and the location of the 11 abrasive notches analysed (red points).
294x128mm (300 x 300 DPI)

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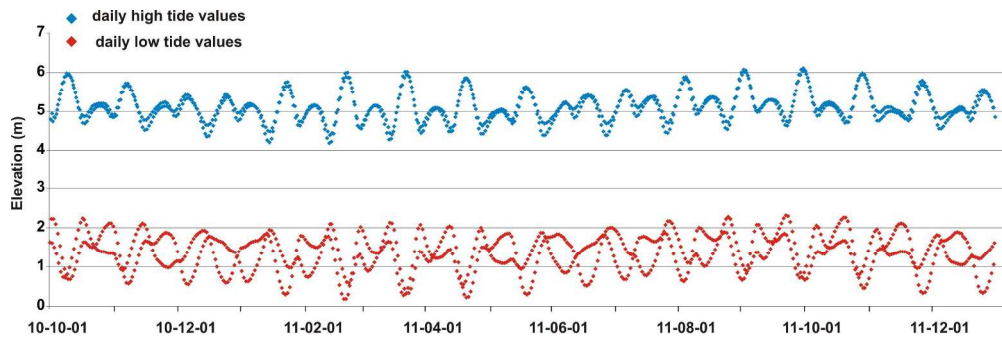


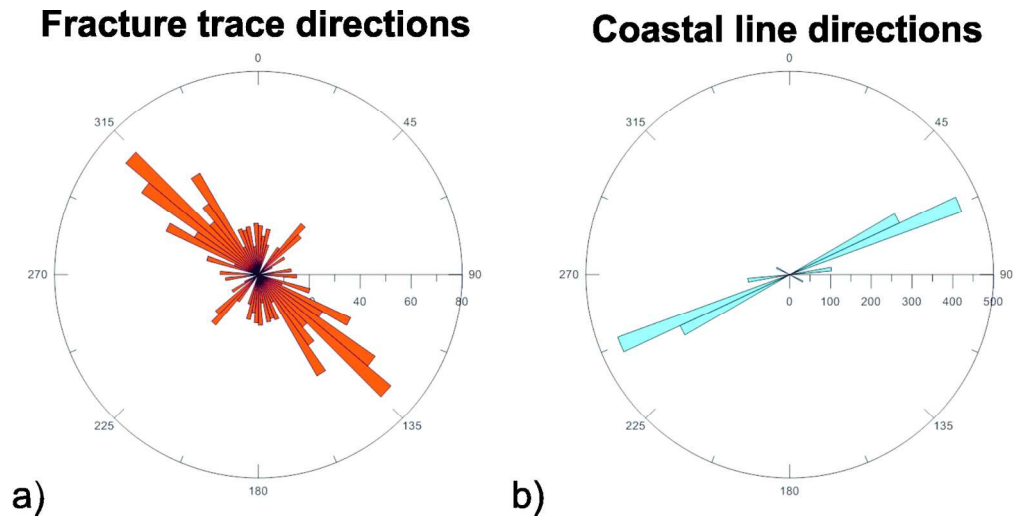
Fig. 2. Tidal trends during the 15 months from 10/10 to 12/11 (http://www.hidro.gov.ar/Oceanografia/Tmareas/Form_Tmareas.asp). The lunar and solar cycles produces in this area a mixed tidal regime yielding unequal low tide and high tide (Max high tide; Max low tide; min high tide and min low tide). Elevation data are in reference to the reduction plane (theoretical plane located under the mean sea level in order to have only positive tidal values in the tables) located 3.20 m below the medium level.

279x90mm (300 x 300 DPI)

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Fig. 3 Examples of the observed notches
193x218mm (300 x 300 DPI)



24 Fig. 4 Analysis of azimuth distribution of fractures (a) and coastal line (b) directions
25 Notice how the predominant fracture direction is approximately at a right angle to the coastline.
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27 154x77mm (300 x 300 DPI)

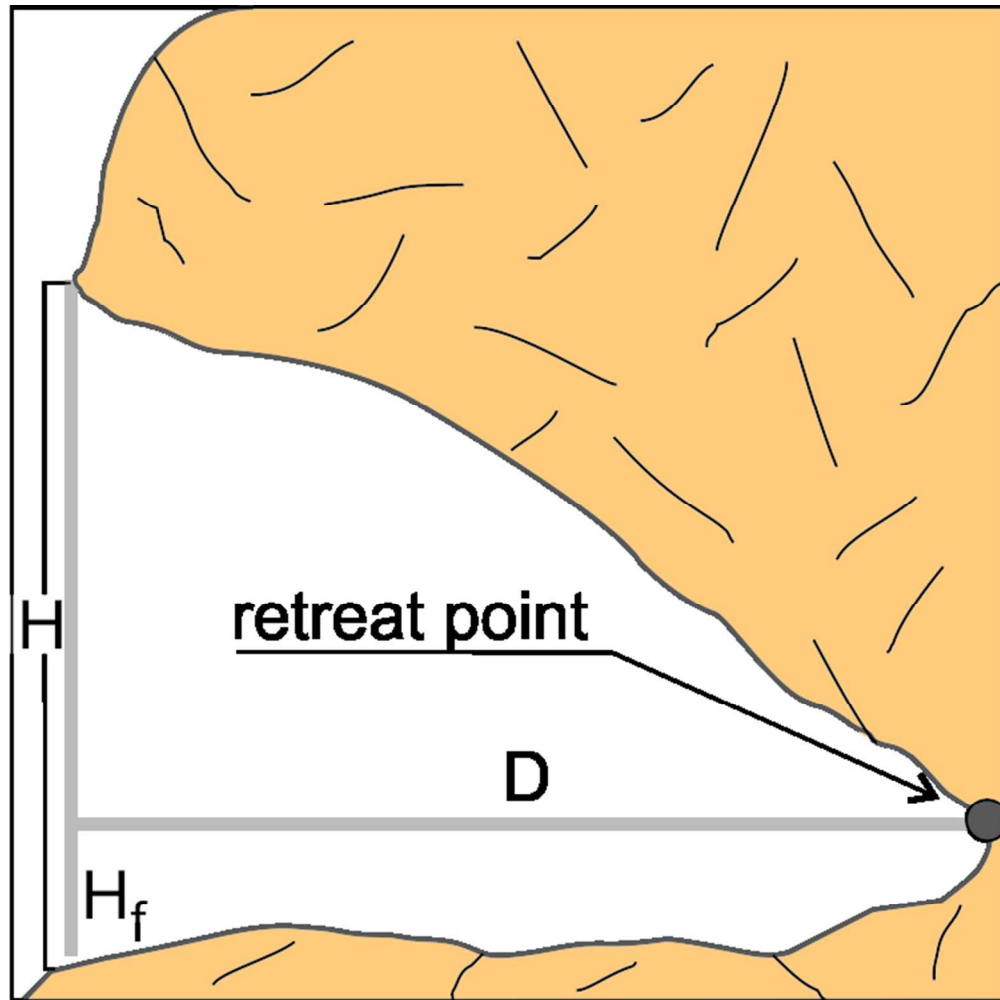


Fig.5 Measured parameters: total height (H), distance of retreat point from cliff face (D), height of retreat point from the floor (H_f), according to Pirazzoli (1986)
63x63mm (300 x 300 DPI)

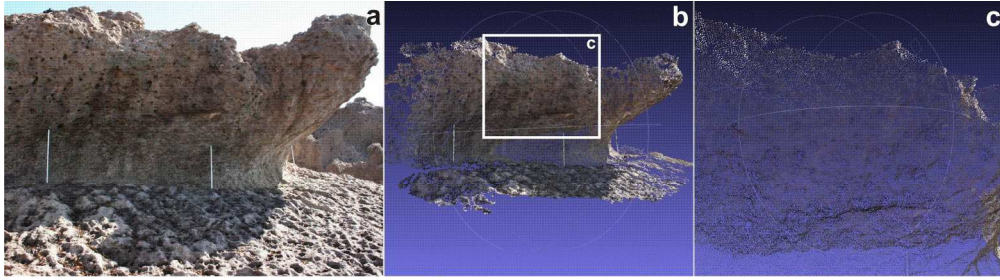


Fig.6 a): Notch1 photography, b): 3D model displayed by MeshLab software, c): zoomed in area on the point cloud.
202x55mm (300 x 300 DPI)

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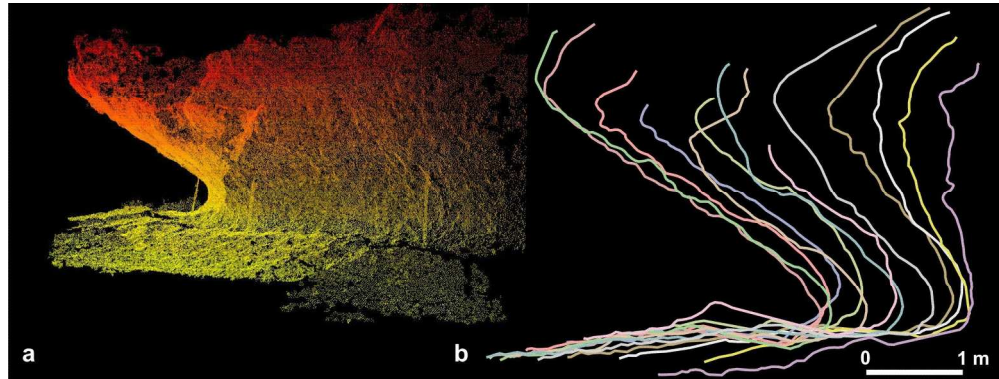


Fig.7 Variability of notches shape, shown through the 3D analysis modelling, a) 3D view of the point cloud,
b) the 13 derived profiles.
204x77mm (300 x 300 DPI)

Peer Review

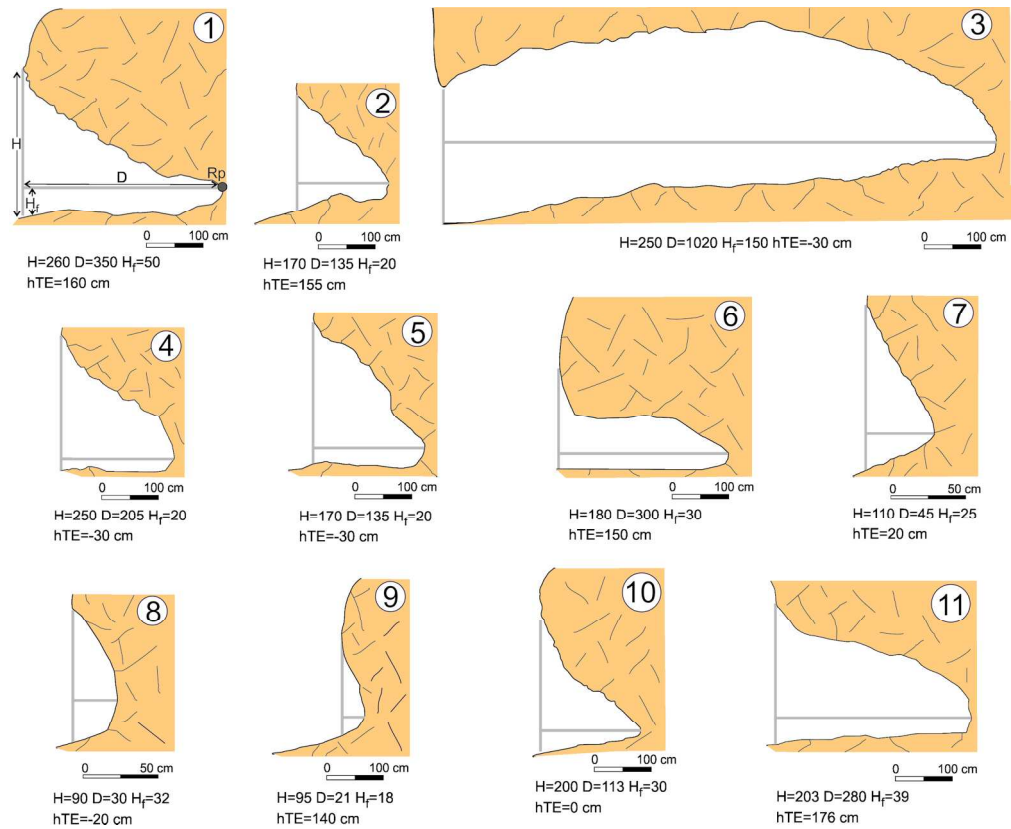
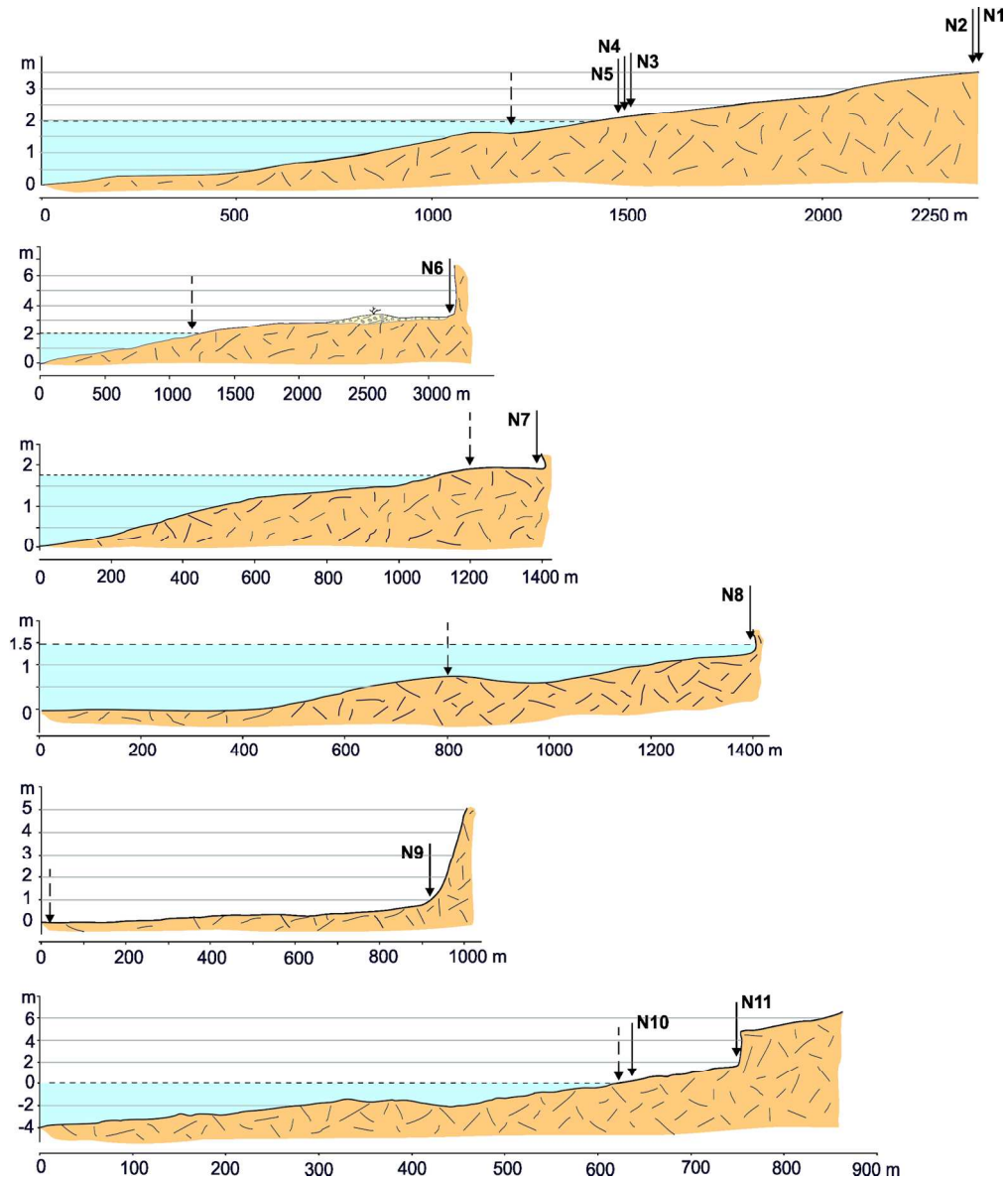


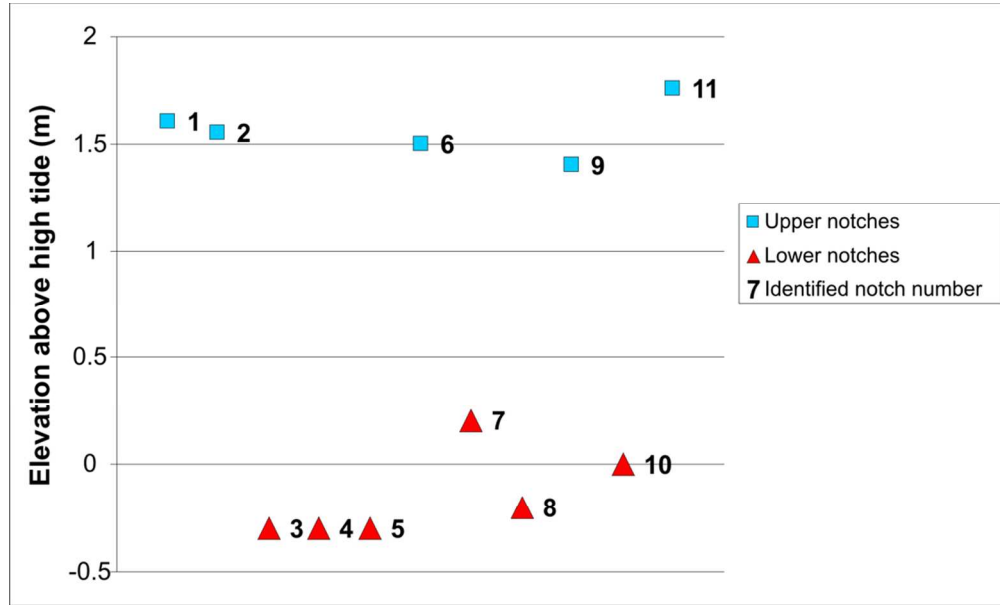
Fig. 8. Morphometric characteristics of the 11 notches analyzed Rp = retreat point; H_f = elevation of retreat point above the floor; H = notch height; D = notch depth; hTE = elevation of retreat point above high tide.

All data are reported in cm and are derived from field measurements.

187x152mm (300 x 300 DPI)

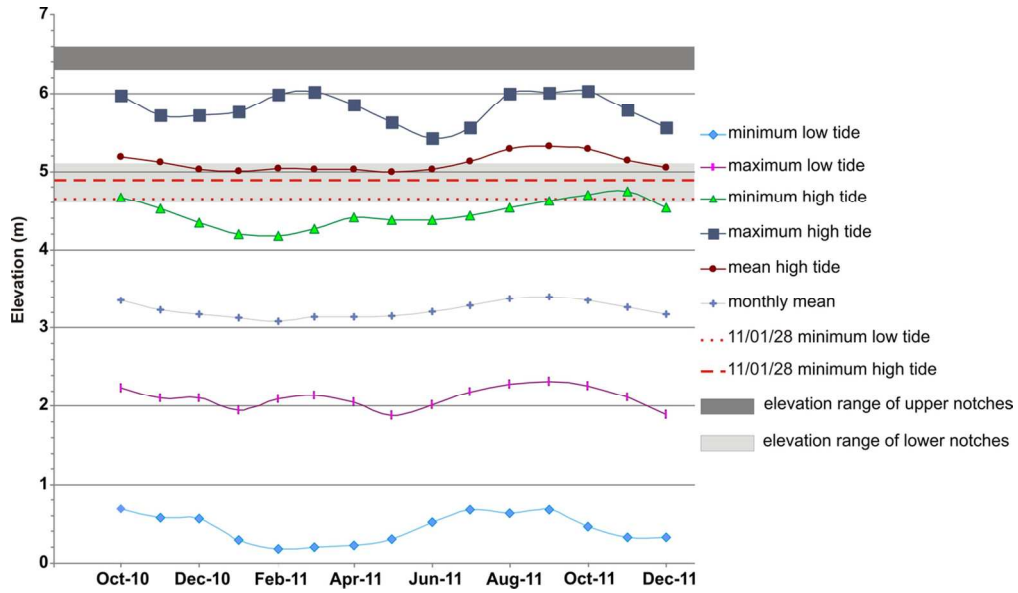


Topographic profiles in which notch position (black arrow) is compared with morphological evidence of high tide (black dashed arrow) and high tide elevation derived from tide tables, calculated for the 2011/01/28 in Puerto Deseado (black dashed line).
 160x188mm (300 x 300 DPI)



Distribution of the retreat point elevations above high tide. In spite the low number of features they show two well separated clusters with a very low variability.
102x61mm (300 x 300 DPI)

Review



Relation between retreat point elevation and tide level derived from the current tide tables of Puerto Deseado. Elevation data are in reference to the reduction plane (theoretical plane located under the mean sea level in order to have only positive tidal values in the tables) located 3.20 m below the medium level.
 117x67mm (300 x 300 DPI)

Review