

**Title: Drimolen crania indicate contemporaneity of *Australopithecus*,  
*Paranthropus* and early *Homo erectus* in S. Africa**

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**Understanding the extinction of *Australopithecus* and origin of *Paranthropus* and *Homo* in South Africa has been hampered by the perceived complex geological context of hominin fossils, poor chronological resolution, and a lack of well-preserved early *Homo* specimens. Here we describe, date and contextualize the discovery of two hominin crania from Drimolen Main Quarry in South Africa. At ~2.04-1.95 Ma, DNH 152 represents the earliest definitive occurrence of *Paranthropus robustus*, and DNH 134 represents the earliest occurrence of a cranium with clear affinities to *Homo erectus*. These crania also show that *Homo*, *Paranthropus* and *Australopithecus* were contemporaneous with at ~2 Ma. This high taxonomic diversity is also reflected in non-hominin species and provides evidence of endemic evolution and dispersal during a period of climatic variability**

**One Sentence Summary:** Multiple hominin genera, including early *Homo erectus*, were present in South Africa at 2 Ma.

**Main Text:**

Remnants of infilled and eroded cave systems (palaeocaves) formed within the Malmani dolomite in north-eastern South Africa have yielded one of the richest early hominin records in the world, including *Australopithecus africanus*, *A. sediba*, *Paranthropus robustus* and early *Homo* (1-6). Most of these hominin species are endemic to southern Africa, with only the early *Homo* material being potentially conspecific with hominins in eastern Africa (1-4), where the origins of *Homo* have been suggested to occur much earlier (7). However, the fragmented nature of most South African early *Homo* specimens (e.g., StW 53; SK 15; SK 847) makes the taxonomy of this group and its relationship to eastern African *Homo* uncertain, with many fossil specimens assigned to multiple species or genera (4). An alternative hypothesis is that *Australopithecus sediba* could be the ancestor of South African *Homo* (8), despite being known from only ~2 Ma (9).

The majority of Early Pleistocene hominin specimens from South Africa come from the Sterkfontein, Swartkrans and Kromdraai B palaeocaves, located within 3 km of each

other in the Blaauwbank Stream Valley, ~40 km north-west of Johannesburg in Gauteng Province (3-6). Prior to 1992, the only other early hominin fossils known from South Africa were the 40 specimens from the Makapansgat Limeworks (~260 km to the north), and the single specimen from the Buxton-Norlim Limeworks (Taung Child; 360 km to the southwest) (6). Moreover, all the discoveries of new early hominin fossils in the last 30 years have come from an area of karst roughly 40 km by 12 km (Fig. 1); referred to here as the 'Gauteng Malmani'. The pattern of hominin evolution in South Africa prior to ~1.1 Ma is thus biased geographically due to the limited extent of the Malmani dolomite karst from which all the fossils derive (5, 6). Correlation of the South African hominin record to that in eastern Africa has been limited by the perceived stratigraphic complexity of many of the South African sites, as well as the historic difficulty in dating palaeokarst due to a lack of suitable material for radiometric dating (5-6). Until recently most dates for South African sites were based on biostratigraphic correlation with the better-dated eastern African sites some 3000 to 4000 km away (5-6).

The stratigraphic sequences at most South African palaeocave sites have been defined based on a lithostratigraphic approach in which breccia deposits and siltstone deposits have been classified as sequential numbered Members, thought to represent different temporal phases of deposition (11-14). In some cases, where stratigraphic contacts can be identified, older lithofacies having clearly collapsed, subsided, or been eroded by secondary cave formation processes before later lithofacies infilled the resulting space (15). This often leads to complexities such as deposits formed by reworking of older units, and thus mixing of fossil material (5, 15-16). In other cases, where stratigraphic links are obscured, the relationship of different lithologies is less certain and the depth that deposits have been accumulated within the cave has often been used to suggest this also represents temporal depth, even when a continuous depositional sequence cannot be identified (6, 13). However, caves do not always conform to the law of superposition and deposits can be inverted, have reworking or intrusive events (5, 16-18). At other sites, where mining or erosion has not obscured stratigraphic relationships it is clear that the different lithofacies represent different depositional processes happening in separate parts of the cave simultaneously, with grading between the different defined Members (19-20). The Members thus bear little relation to chronostratigraphy and there is profound lateral

variation in lithology within what should strictly be defined as a single Member. Misinterpreting these complexities has led to much confusion when defining the stratigraphy of the sites, which often requires a well-resolved chronology for robust interpretation (5-6, 15,19).

With the advent of uranium-lead (U-Pb) geochronology as a method for dating Pliocene and Pleistocene speleothems it became possible to date the flowstones that underlie and cap fossil bearing deposits, creating datable Flowstone Bound Units (FBUs) (21-22). Dating of flowstones across the Gauteng Malmani revealed their contemporaneous formation in multiple caves between ~3.2 and ~1.3 Ma (21). As such, flowstones can be used to derive a regional chronology in the same way as volcanic tuffs in eastern Africa (21). However, due to mining and/or surface erosion at some sites capping and underlying flowstones have been removed (19), and the dating of flowstones alone may only provide broad age ranges for the associated fossil-bearing cave sediments between them (21). It is therefore critical to combine U-Pb dating with other complementary methods such as Uranium-Series Electron Spin Resonance (US-ESR) dating and palaeomagnetism (5, 6, 15, 19). As complex depositional situations can occur in caves, such as speleothem false floors, it is also important to show the nature of the contact between sediments and dated speleothem using micromorphology.

Recent geochronological work in the Gauteng Malmani suggests a transition occurs between ~2.3 and ~1.8 Ma from supposedly older sites containing *Australopithecus* (Malapa, Sterkfontein Member 4) to supposedly younger sites containing *Paranthropus* and *Homo*, together with the first bone and stone tools (Gondolin, Kromdraai B, Sterkfontein Member 5, Swartkrans Member 1) (5-6, 9, 15, 21, 23-25). At the same time there is a turnover in other fauna as South African environments became more arid (26-29). However, coarse chronological resolution and imprecise provenance of historically collected fossils (5) have limited interpretations of how and when faunal communities changed and whether *Paranthropus* and/or supposed early *Homo* fossils directly relate to earlier *Australopithecus* species, or dispersed into the region and ultimately replaced them.

Work on sites outside the Blaauwbank stream valley has expanded our understanding of the South African record and has revealed that not all palaeocave sites have complex multi-generational phases of karstification and infill (9, 19) as documented

at sites like Sterkfontein, Swartkrans and Kromdraai (5-6, 15, 17). Discovered in 1992, the Drimolen palaeocave complex (Fig. 1) is one such site (30). Drimolen has yielded over 155 hominin specimens (1, 31) together with significant collections of other fauna (32), bone tools (25), and a small assemblage of Mode 1 stone tools (24). The younger part of the Drimolen system, known as Drimolen Main Quarry (DMQ), is best known for the 1994 discovery of the DNH 7 cranium, the most-complete *P. robustus* skull found to date (1). Much of the rest of the DMQ hominin material consists of isolated teeth (31). However, in 2015 and 2018 two new hominin crania were discovered representing *Homo* (DNH 134) and *Paranthropus* (DNH 152) respectively. These recent fossil finds, together with a well-resolved chronology at DMQ, now make it possible to address in greater detail the complex period of change in hominin evolution that occurred around 2 Ma and sets a standard for dating fossil-bearing palaeokarst.

### **The DNH 134 *Homo aff. erectus* cranium**

DNH 134 comprises a partial neurocranium (Fig. 2) preserving most of the occipital squama, parietals, and frontal squama with no evidence of plastic deformation. The cranial sutures are patent and at an early stage of fusion, indicating that the specimen is a juvenile. The metopic suture is fused externally and the anterior and posterior fontanelles are absent, indicating an age at death greater than 12-36 months according to modern human standards (33-34). The parietals exhibit two tables of bone separated by diploë, indicating that DNH 134 was ontogenetically older than the Mojokerto juvenile (35-36 but see 37). To estimate cranial capacity we created a partial virtual endocast and used multivariate statistics based on three-dimensional landmark data of a reference sample (Fig. 3; see materials and methods). The endocranial volume as predicted via linear regression is 538 cc with a 95% single prediction band from 514 to 564 cc. Estimates based on multiple thin-plate spline reconstructions have a larger range but are consistent with this estimate (484-593 cc). Thus, estimated brain size in the juvenile DNH 134 overlaps with the high end of the range of adult *Australopithecus* and *Paranthropus* but exceeds the brain sizes (275-410cc) of juvenile *Australopithecus* (2, 38-39). While extrapolation of an adult cranial capacity is not straightforward, it is clear that at this ontogenetic stage, DNH 134 has not reached adult size but possesses a cranial capacity at the lower range of adult variation of the population

from which it was drawn. Assuming an age at death between two and three years, DNH 134 could have reached a cranial capacity between 588-661 cc or 551-577 cc according to a human or a chimpanzee growth model, respectively.

The specimen preserves characters that align it morphologically with *H. erectus sensu lato* (including *H. ergaster*): its profile is “teardrop” shaped in superior view, its squamosal suture is nearly straight, sagittal keeling is present on the frontal and parietals, the cranial vault is long and low with strong sagittal occipital curvature and lambdoidal flattening, and although the anterior aspect of the foramen magnum is missing it is evident that a basion – bregma chord would have been short. These traits together distinguish DNH 134 from *A. africanus*, *P. robustus* (as preserved in DNH 7), *H. habilis*, *H. rudolfensis* and *H. naledi* (40). Individually, none of these traits is fully diagnostic of *H. erectus s.l.*, which is morphologically variable across time and space (41), yet collectively they strongly suggest an affinity with that species. Indeed, DNH 134 is strikingly similar to the Mojokerto *H. erectus* cranium in overall cranial shape (Fig. 4).

### **The DNH 152 *Paranthropus robustus* cranium**

DNH 152 is a partial cranium preserving much of the left side of the parietal and frontal bones, a portion of the occipital, the right temporal, the lateral margin of the right orbit, as well as four teeth: left and right maxillary first and maxillary second molars (Fig. 5). The right dental elements are *in situ* within a portion of the maxilla while the left elements are isolated. The specimen preserves a number of characters that align it taxonomically with *P. robustus*. The supraglenoid gutter is partially preserved and would have been wide. The mastoid process is inflated lateral to the supramastoid crest, from which it is separated by a broad shallow groove. The external auditory meatus is large and nearly circular (11.8 mm by 11.5 mm), positioned lateral to the tip of the mastoid process, and nearly at the level of the suprameatal crest. The post-glenoid process is small and fused to the tympanic. Although the tips of the articular tubercle and entoglenoid process are missing, it is evident that the articular eminence was wide and that the glenoid fossa was deep. The digastric groove takes the form of a narrow notch. The superior temporal line on the frontal bone is a well-demarcated ridge, suggesting that a frontal trigon would have been present, but this cannot be directly observed. The sagittal crest bifurcates superior to

lambda leaving a bare area on the occipital, and the temporal lines meet the nuchal line in the lateral third of its extent forming a short, partial compound temporonuchal crest. There are extensive striations and beveling on the inferior aspect of the left parietal bone, indicating that the overlap between the temporal and parietal bones at the squamosal suture was extensive. The inferior orbital margin is rounded laterally. Both molars evince a quadrangular occlusal outline with mesiobuccal extension, a deep and narrow central fossa and longitudinal fissure, and a thick distal marginal ridge typical of *P. robustus*. The left M<sup>1</sup> is fractured and shows ‘hyper-thick’ enamel. Based on these characteristics the cranium has been assigned to *P. robustus*.

### **Morphology of the Drimolen Main Quarry (DMQ) Palaeocavern**

The Drimolen palaeocave system consists of at least two temporally and spatially distinct cave infills: the ~2.61 Ma Drimolen Makondo (DMK) deposit (19), and the hominin-bearing DMQ (27-29) (Fig. 1C). DMQ is a large palaeocavern (~20 m x 15 m) formed in one of the highest current exposures of the Gauteng Malmani (~1545 m amsl). Based on current topography, the palaeocavern had a very small watershed and acted as a vertical pothole sink for groundwater. The nearby Wonder Cave (aka: van Wyk’s Main Cave) (42) provides a good modern analogy. In contrast, cave sites in the Blaauwbank Valley, such as Sterkfontein and Plovers Lake act as valley bottom ‘collecteur’ caves, estavelles or exsurgences where underground lakes occur either intermittently or permanently (43). While the landscape in the Gauteng Malmani has been altered over the last few million years due to erosion (9), this would not have significantly changed the watershed of the Drimolen palaeocave due to its location near the top of a hill, rather than within an actively incised valley. Moreover, only ~10m of erosion is estimated per million years on the hill behind DMQ (43). If washed in, the sediments, fossils, and archaeology deposited in DMQ would have originated from a restricted landscape around the cave.

Today, DMQ is a roughly subcylindrical karstic depression with vertical sides, about 5-8 m deep. Most of the ceiling and the upper parts of the DMQ cavern infill have been lost to erosion, breakdown and dissolution (Fig. 1, 6). The outline of the current quarry was created through speleothem (lime) mining in the late 19<sup>th</sup> century. The pattern of mining indicates that speleothem deposition originated on the eastern side of the

palaeocavern. Water forming these speleothems flowed down-slope along the bedding of the Malmani dolomite (to the north-west) to form a thinning flowstone floor on the south-western side of the cavern. The combination of natural erosion/dissolution and anthropogenic mining has produced a range of stratigraphic profiles that exposes the entire formational history of the cavern (Fig. 6-10).

The western part of DMQ consists of a 10m wide belt of *in situ* calcified palaeocave sediments adhering to the western dolomite wall of the DMQ palaeocavern and deposited over a remnant of basal flowstone. These *in situ* deposits consist of calcified sediment pinnacles up to 3 m high, shaped by subcutaneous secondary karstic dissolution. The spaces between these pinnacles (makondos) (19) are filled by soft sediment that represents the *in situ* decalcified equivalent of the sediment in the pinnacles and articulated bone can be found embedded across the contact of both mediums (19). Both DNH 134 and 152 were recovered partly from decalcified and partly from lightly calcified breccia, and in close contact to solid breccia. Directly to the east of these *in situ* deposits, are a mixture of collapsed and decalcifying palaeocave deposits and miner's rubble that were the focus of excavations between 1992 and 2016 (30). This 'Central Excavation Area' has yielded most of the fossil material, but it is mostly *ex situ*. This includes the DNH 7 *P. robustus* cranium (1) that comes from a large block (Eurydice Block: Fig. 6) in the centre. Fossil material in this collapse zone, and in the centre of makondo features (19), often shows poorer preservation, which is likely why the majority of hominin fossils recovered to date consist of isolated teeth (31). The collapse of this breccia was caused by the undermining of the *in situ* deposits during lime-mining forming the Inner Cave (Fig. 7). While secondary cave formation has also occurred at DMQ it is limited to the formation of Warthog Cave at the contact between the palaeocavern infill and the southern dolomite wall of the palaeocavern (Fig.7). Warthog Cave has only slightly eroded into the palaeocave fill and the only fossil to have been recovered from the fill is an isolated *Paranthropus* molar (DNH 122) that lay very close to the contact with the palaeocave deposits. There is no evidence that this cave has affected the palaeocave deposits via other mechanisms such as collapse.

### **Stratigraphy of the Drimolen Main Quarry (DMQ) Palaeocavern**



Here we outline the geochronology and stratigraphy of the DMQ palaeocavern as recorded in a number of key stratigraphic sections along the mined and excavated exposures of the *in situ* western wall palaeocave deposits (see also Supplementary Video 1). The various lithofacies described in Table S1 have been identified by stratigraphic analysis and micromorphology (44). The deposits represent a process of continuous accumulation, erosion and reworking of a single, large clast-supported talus cone breccia (CSB) that formed on a basal flowstone (BFS) beneath a vertical entrance. This talus cone was then subject to lateral winnowing during flooding to sequentially create matrix-supported breccia (MSB) and distal gravel, sandstone and siltstone (GSS) (Fig. 11, 12; Table S1). Micromorphological observations (Table S1; Fig. 11A-B) indicate the in-washing of sediment derived from the erosion of colluvial soils previously developed outside the cave. A lengthy evolution for these soils is indicated by the intense weathering of minerals and rocks (Fig. 11B). The granular microstructure of the breccia sediment mass is generally rather loose (Fig. 11A), mostly in the upper part of the breccia cone, suggesting fast deposition and leaching of the fine particles, followed by rapid cementation of the sediments. Calcite is the most common cement for both breccias and sandstone/siltstones (Fig. 11D), with the most frequent crystalline pattern being mosaic calcite with anhedral crystals of variable size that cements the whole mass. Fragments and splinters of compact or cancellous bone, as well as microfauna, are often present throughout the breccia (Fig. 11C). The breccia also contains well-preserved fragments of vegetal tissue, which exhibit cellular patterns typical of large-size monocotyledon taxa, and were likely washed in (Fig. 11C). The shape of voids in the breccia are consistent with formation by roots and/or burrowers (Fig. S11C). These occur in samples from about 30 cm above the basal flowstone indicating the early opening of a reasonably wide entrance, and thus indicating that no upper cavern existed as previously suggested (30).

**Warthog Cave Section** is the most southerly and deepest exposure of the DMQ palaeocavern fill (Fig. 7). The base of the 3m deep section consists of a 50 cm thick flowstone speleothem (which was thicker before mining) that contains no significant detrital material and formed before the cavern had an opening to the surface. This basal flowstone (BFS; -6.40 m below datum) has been U-Pb dated to  $2.673 \pm 0.103$  Ma (DN39A; 21). The normal polarity of the flowstone limits its formation to older than the Gauss-

Matuyama Boundary at 2.61 Ma (45), setting a lower age limit on the DMQ deposits. There is a sharp contact with an overlying fossil-bearing, clast-supported breccia (CSB) that formed the extreme southern toe of the talus cone down the westerly dipping flowstone and represents the oldest fossil bearing deposits at the site (Lower Cave Breccia). This unit is not noted elsewhere, has not been excavated, and is overlain by a series of well-stratified GSS deposits (Warthog GSS; Table S1). WGSS represents winnowing of fine-grained material from near-entrance talus deposits to the southern edge of the cavern during floods. WGSS was sampled for palaeomagnetic analysis and recorded a reversed polarity (Fig. 13; Table 2) consistent with being deposited between BFS and WCFS, between 2.61 and 1.95 Ma. WGSS filled the southern part of the palaeocavern to the low stepped roof that now forms the top of Warthog Cave. The top of WGSS has been eroded and capped by a 15 cm thick flowstone (Warthog Cave Flowstone; WCFS; -3.90 below datum), which infilled an erosional channel between the palaeocave deposits and the western dolomite wall of the palaeocavern. The flowstone dates to  $1.789 \pm 0.104$  Ma by U-Pb (DN09; 21) and recorded a normal magnetic polarity (Fig. 13; Table 2) consistent with deposition during the Olduvai SubChron (1.95-1.78 Ma), and indicating its formation 170-60 ka after the deposition of the WGSS deposits on which it lies (45).

**Jangi Buttress** occurs just to the north of the Warthog Cave Section at the southwest edge of the Central Excavation Area, in the centre of the site (Fig. 8). The Jangi Buttress comprises a 3 m deep outcrop of clast-supported breccia (CSB; Table S1) representing a talus cone breccia formed from a vertical entrance and accumulated against the western wall of the palaeocavern overlain by matrix-supported breccia (MSB; Table S1). The breccia consists of large angular to sub-angular dolomite and chert blocks (up to ~40 cm) formed by entrance and roof collapse, with small pockets of fine-grained in-washed sediment, often with dense microfossils, occurring between the blocks. A continuous outcrop of CSB extends from Jangi Buttress west to the Western Wall and then north to the Italian Job Pinnacle (see below) (Fig. 6). Adhering to the Jangi Buttress on its northern side (western wall of Central Excavation Area) are decalcified remnants of CSB. Excavation of this decalcified material mimics the nature of the breccia with pockets of yellowish red micromammal rich sediment and ghost rock nodules consisting of the insoluble fraction of decalcified dolomite boulders and cobbles. The DNH 134 *Homo*

cranium was recovered as a series of individual pieces at ~-5.31 m below datum (1.1 m above BFS) towards the base of the pinnacle (Fig. 8) during excavations in 2008, 2015, 2016 and 2019 (from the single piece from 2008 was not recognized as hominin until more of the cranium was recovered in 2015). The cranial pieces were partly recovered from lightly decalcified CSB and from decalcified sediment and collapse (since 2008 excavations) next to the pinnacle. These deposits also yielded adult *Paranthropus* teeth and bone tools. MSB and CSB of the Jangi Pinnacle is equivalent to the GSS deposits of the Warthog Cave section and represents the talus cone from which the WGSS deposits were winnowed. A US-ESR age from a bovid tooth next to the cranium gave an age of  $2.041 \pm 0.240$  Ma (see below; Fig. 8), further confirming this association, and shows the WGSS and Jangi Buttress CSB sediments were deposited ~600-280 ka after BFS formed at >2.61 Ma.

**The Italian Job Pinnacle Section** is located just to the northwest of the Central Excavation Area (Fig. 9) and occurs stratigraphically higher than the Jangi Pinnacle, with which it is connected by *in situ* breccia that also connects both sections to the West Wall of the palaeocavern. The Italian Job Pinnacle consists of an intermediate facies of MSB (Table S1) that represents a vertical and lateral transition from CSB at the core of the central talus cone to MSB as fine sediment begins to dominate over large clast deposition during the vertical and westward expansion of the central debris cone. It also represents a lateral transition from MSB of the upper part of the central debris pile to GSS deposits that occur on the northern side of the Italian Job Pinnacle itself and in the Walls of Jericho Pinnacle on the northern edge of the palaeocavern (see below). This transition represents winnowing of fine-grained sediment from the central talus cone to the northern corner of the cavern during floods. A thin flowstone (Walls of Jericho Flowstone; WOJFS) occurs at -0.97 m below datum in the Italian Job Pinnacle and continues into and through the adjacent Walls of Jericho Pinnacle, suggesting a slight hiatus in deposition. US-ESR dating of a bovid tooth from ~18 cm below the flowstone at -1.15 m below datum produced an age of  $1.965 \pm 0.147$  Ma (see below). The DNH 152 *Paranthropus robustus* cranium was recovered from the very base of the current excavated exposures of the Italian Job at a height of -3.15m below datum and ~2.18 m below the WOJFS.

**The Walls of Jericho Pinnacle Section** is the most northerly exposure of the DMQ palaeocavern infill (Fig. 10). Most of the section comprises GSS, representing fine-grained sediments winnowed by medium-energy flow (Fig. 11E-F) from the central debris pile to the south. On the northern side of the pinnacle a steeply dipping flowstone is assumed to be equivalent to the basal flowstone in the Warthog Cave section (2.78-2.61 Ma) (DN39A; 21). As in the Warthog Cave section, there is a sharp contact between this basal flowstone and the overlying sediments. At -0.78m below datum the WOJFS occurs (~2cm thick) that has been U-Pb dated to  $1.962 \pm 0.107$  Ma (DN26) (21), consistent with the US-ESR age just below this flowstone in the Italian Job Pinnacle. Micromorphological analysis (Fig. 12) confirms this flowstone is not intrusive and formed during a hiatus in the deposition of GSS. GSS deposits below the WOJFS are well-laminated, whereas those above are more coarsely layered, with thin intercalated flowstones and silt crusts suggesting alternating hiatuses in clastic deposition and pooling of water (Fig. 11G-I). Palaeomagnetic analysis indicates a change from reversed to normal polarity up through the section with intermediate polarity occurring in the WOJFS as well as sediments above and below it, further confirming it formed during the deposition of the sediment sequence (see below). Based on the U-Pb and US-ESR ages this can be correlated to the reversal at the base of the Olduvai SubChron at ~1.95 Ma (46).

### Geochronology

US-ESR analysis was undertaken on an indeterminate medium-sized alcelaphin right maxillary third molar (right M3; DMQ-2) recovered from the lightly decalcified breccia of the Jangi Buttress in direct association with the DNH 134 cranium (Fig. 7); and, on another partial bovid tooth (DMQ-3) encased in breccia from the southern side of the Italian Job Pinnacle, 2 m above DNH 152 and around 20 cm below the WOJFS (Fig. 9). The Dose equivalents for DMQ-2 and DMQ-3 were estimated using the peak-to-peak  $T_1-B_2$  method at  $1814 \pm 57$  and  $2414 \pm 76$  respectively ( $2\sigma$  error) (Table S2) including an unstable radical component (NOCOR) of 21% and 16%, respectively (47-49). When integrated into the US-ESR dating modelling described by Shao et al. (50), the ages of DMQ-2 and DMQ-3 are estimated to be  $2.041 \pm 0.240$  Ma and  $1.965 \pm 0.147$  Ma respectively ( $1\sigma$  error) (Table 1, S2). Both samples did not show any ratios above secular

equilibrium, although the dental tissues were not extensively mapped. Nonetheless, the U-diffusion in DMQ-2 and DMQ-3 fits the open-system model, and the teeth did not exhibit obvious complex compound uranium diffusion episodes. However, the isotopic ratios between the enamel and dentine remain different, most likely indicating a more recent incorporation (uptake) of uranium in the dentine. With a U-uptake history model close to linear in most dental tissues (Table 1) as well as rather homogenous ratio over the analysed area, we were able to assume equilibrium in the uranium decay chain after  $^{230}\text{Th}$  (e.g.,  $^{210}\text{Pb}/^{230}\text{Th}=1$ ).

Palaeomagnetic analysis was performed on the GSS and U-Pb dated flowstones from the Warthog Cave and Walls of Jericho Sections (Table 2; Fig. 7, 10). Natural Remanent Magnetisation (NRM) intensities ranged from 8.94 to 0.57  $\text{Am}^2/\text{kg}^1$  with a mean  $\chi_{\text{LF}}$  (non-speleothem samples) of  $186 \times 10^{-5}$  SI indicating sufficient ferromagnetic material for palaeomagnetic analysis. Mineral magnetic measurements (Fig. 13G-I) show that both low coercivity magnetite and maghemite occur in superparamagnetic (SP), stable single domain (SSD), and vortex state (formally referred to as pseudo-single domain) (51) grain size ranges. Pigmentary hematite is also likely present as indicated by a low contribution antiferromagnetic component observed in backfield unmixing curves (Fig. 13G). This does not contribute to the NRM. A significant proportion of SP to viscous single domain (vSD) boundary grains are shown by high frequency dependence of magnetic susceptibility ( $\chi_{\text{FD}}\%$ ; mean 11.75% for non-speleothem samples) and these are susceptible to more recent viscous re-magnetisation. First Order Reversal Curves (FORCs) show a predominance of single domain signals, which are ideal for palaeomagnetism, and are highlighted by closed concentric contours along the central ridge of the FORC diagram (Fig. 13H) (52). Given that some of the central ridge coercivity distribution intersects with the Bu axis the SSD magnetisations are likely mixed with those of lower coercivity SP/vSD boundary grains (53) in line with  $\chi_{\text{FD}}\%$  results. Such samples require caution when undertaking  $\text{AF}_D$  cleaning to ensure remanences associated with SSD particles are properly isolated from SP/vSD boundary overprints prior to their removal. FORCs also derive some influence from larger vortex state particles as shown by a spread of remanence away from the central ridge and weak lobe features, while there is no evidence for multi domain (MD) grains, which would be reflected by a greater remanence spread along the Bu axis (51).

Final palaeomagnetic data for the 13 block samples are presented in Table 2 and Fig. 13. Viscous overprints associated with SP/vSD boundary grains were removed via  $AF_D$  typically between 12–15 mT or with  $TH_D$  at temperatures of 250–325 °C to reveal a single stable component of magnetisation comprised of low MADs (<10; Fig. 13A-F). This ChRM signal is removed via  $TH_D$  between 540–580 °C (and 5–40mT via  $AF_D$ ), suggesting detrital magnetite ( $T_c = \sim 585^\circ\text{C}$ ) as the main remanence carrier, corroborating  $T_c$  estimates of 561–586 °C derived from  $M/T$  heating curves (Fig. 13I). In isolated samples the ChRM wasn't removed until  $\sim 610$  °C suggesting some minor maghemite contribution. Taken together, the overall stability of magnetisation to high temperature along with the occurrence of reversed polarities and directional consistency among different demagnetisation strategies indicates a primary remanence formed within SSD (to vortex state) magnetite and maghemite around the time of sediment deposition in the cave (i.e. depositional or post-depositional remanent magnetisation; DRM or pDRM). The occurrence of microlayers of flowstone within the sediments likely aided the quick lock in of magnetic remanence and reduced any effects of post-depositional remanent magnetisation. Final directions (Figs 7, 10, 13, 14; Table 2) indicate a series of polarity changes throughout the sequence. Normal polarity is observed in the basal and capping flowstones of the Warthog Cave section and the upper 60cm of the GSS of the Walls of Jericho Section. Reversed polarity is noted in the GSS of the Warthog Cave Section and base of the Walls of Jericho Pinnacle. The middle part of the GSS Facies (-1.55 to -0.59 m below datum) in the Walls of Jericho Section, as well as the WOJFS itself, record intermediate magnetic polarity that represents true intermediate geomagnetic behaviour occurring during a magnetic reversal. It is extremely rare to find evidence for such reversals in terrestrial sedimentary sequences and indicates this part of the sequence likely formed over a few to several 1000 years during the reversal itself.

### **Chronostratigraphy and Flowstone Bound Units**

Despite the perceived stratigraphic complexity of Gauteng Malmani palaeocaves, DMQ has a simple depositional history. Evidence suggests the site was a water sink (pothole) that infilled over a short timeframe and there is little evidence for secondary cave formation, natural reworking of fossils, or infills of significantly different ages. Putting all

the geochronological data and stratigraphic information together, the DMQ sequence can be divided into two major flowstone-bound units (FBU; Fig. 14; Table 2). The thicker (~5.6 m thick) stratigraphically lower FBU (FBU1) occurs between the  $2.673 \pm 0.103$  Ma (16) Basal Flowstone (-6.40m below datum) that underlies the entire sequence of clastic deposits within the palaeocavern and the  $1.962 \pm 0.107$  Ma Walls of Jericho Flowstone (-0.78 m below datum) that divides the Walls of Jericho and Italian Job Pinnacles (Fig. 12) (21). The thinner (~1.07m), upper FBU (FBU2) formed between the Walls of Jericho Flowstone and the  $1.789 \pm 0.104$  Ma Warthog Cave Flowstone (Table 2) (21). The normal polarity of the Basal Flowstone indicates it formed prior to the Gauss-Matuyama boundary at ~2.61 Ma (45), while the sharp contact with overlying sediments and lack of detrital inclusions confirms the flowstone formed prior to an opening to the surface (as seen in base of Warthog Cave Section).

During the formation of FBU1 a large clast-supported breccia (CSB) was deposited beneath a vertical entrance in the western to central part of the palaeocavern (as seen in Jangi Buttress) and flooding winnowed fine-grained sediments (GSS) against the southern wall of the palaeocavern (as seen in Warthog Cave Section). Between the clasts within CSB fine-grained sediments (equivalent to GSS) and macrofossils occur in small pockets. In this early phase articulated skeletons are often found across more than a single pocket. This indicates deposition of the skeletons before or during the deposition of blocks rather than a mechanism whereby fossils and fine sediment have filtered down through a pre-existing structure of blocks. The reversed polarity of the Warthog Cave GSS and US-ESR age of  $2.041 \pm 0.240$  Ma for the oldest excavated CSB deposits from the Jangi Buttress indicate the oldest sediments, including *Homo aff. erectus* fossil DNH 134, entered the DMQ palaeocavern several 100,000 years after the Basal Flowstone formed, and sometime between ~2.04 and 1.95 Ma, during the Matuyama reversed polarity Chron. Isolated teeth of *Paranthropus robustus* and bone tools have also been recovered from these layers. As the central talus cone continued to be deposited in the central part of the cavern it also expanded against the western dolomite wall of the cave. As the CSB of the talus central talus cone built up it also graded laterally to a more matrix supported breccia (MSB) in the western half of the cavern (as seen in the base of the Italian Job Pinnacle). At this time, when DNH 152 was deposited, the GSS deposits has already filled the cavern to the roof

in its lower southern part and GSS deposits were now being winnowed against the northern wall where the cavern roof was higher (as seen on northern edge of Italian Job and in the Walls of Jericho Pinnacle). Like the top of the Warthog Cave Section, the oldest deposits in the Walls of Jericho Pinnacle record a reversed polarity. The DNH 152 *P. robustus* cranium was deposited during the middle part of FBU2 ~3.2 m above the basal flowstone, 2.2 m above DNH 134, 2.2 m below the WoJFS, and ~1.6 m below the beginning of the magnetic reversal in the Walls of Jericho (45-46) (Fig. 14).

The talus cone continued to form and as sediment was further winnowed into the northern edge of the cave (Italian Job to Walls of Jericho Section) the polarity of the GSS deposits changed from reversed to intermediate polarity, indicating the onset of a magnetic reversal. The US-ESR age of  $1.965 \pm 0.147$  Ma for these deposits suggests this reversal is the onset of the Olduvai Subchron at ~1.95 Ma. This is confirmed by the  $1.962 \pm 0.107$  Ma U-Pb age and intermediate polarity for the thin WoJFS (DN26; 16) that formed during a hiatus (shorter than the few to several 1000 years of a magnetic reversal) in deposition and caps the lower well-laminated GSS deposits in the northern part of the cavern (as shown in the WOJ; Fig. 10, 14). GSS continues to form after the WoJFS and again records intermediate polarity before transitioning to the normal polarity of the Olduvai Subchron just prior to GSS filling up to the roof of this part of the cavern (1.95-1.78 Ma) (46). This indicates that the upper parts of the lower FBU, the WoJFS, and the lower part of the upper FBU were all deposited during the timeframe of the magnetic reversal at the base of the Olduvai Subchron at ~1.95 Ma. Reversals have been estimated to take between 4,000 and 22,000 years to complete (54). Significantly after this event, an erosional channel formed between the talus cone and the western dolomite wall, also eroding the top of the GSS sediments in the Warthog Cave section. This erosional channel was filled with the Warthog Cave Flowstone (WCFS) that has been U-Pb dated to  $1.789 \pm 0.104$  Ma (DN09) (21) but whose age can be refined to between 1.89-1.78 Ma due to its normal polarity and correlation to the Olduvai Subchron (1.95-1.78 Ma). The WCFS caps FBU2, although most of the fossil-bearing sediments in FBU2 only occur in the top ~90cm of the Walls of Jericho and Italian Job Pinnacles (Fig. 14) and are formed immediately after the reversal at ~1.95 Ma. While a younger US-ESR age of  $1.712 \pm 0.538$  Ma has been produced for a tooth from the Eurydice (DNH 7) block that may suggest deposition during FBU2, the *ex situ* nature



of the block and its decalcification make dosimetry estimates more difficult. The Eurydice block consists of CSB and is consistent with deposition in FBU1 and the US-ESR age does overlap with the other ages from FBU1 within error. Moreover, comparisons between the ESR dating of teeth from calcified and decalcified breccia at other sites show that decalcification causes ages that are too young (15). As such, it is critical to collect teeth from *in situ* calcified breccia when undertaking US-ESR analysis at such sites. The lack of any short magnetic reversal events, such as the ~2.07 Ma Huckleberry Ridge event (45) identified at other palaeocaves in the region (9, 15) within deposits older than the reversal at DMQ suggests the deposits all formed post ~2.07 Ma, consistent with the median age of the US-ESR sample from the Jangi Buttress. The vast majority of hominin remains from DMQ, including DNH 134 and DNH 152, thus come from FBU1 maximally between 2.28 and 1.95 Ma, but most likely between ~2.04 and 1.95 Ma.

### **Biogeographic Interpretations of the Drimolen Main Quarry Faunas**

There are several faunal species represented at DMQ (32) that support some level of non-endemic mammal dispersal into South African palaeoecosystems during the early Pleistocene, and/or temporal variation within South African phyletic lineages. The recovery of *Equus* cf. *quagga* ssp is consistent with the deposits having formed after ~2.3 Ma given the first appearance of the genus in eastern Africa at this time (32). *Equus* is first seen in South Africa at Sterkfontein Member 4 prior to ~2.07 Ma (6, 15), as well as at Malapa just after ~2 Ma (9). This indicates the relatively rapid expansion of this genus into the southern part of the continent, perhaps coinciding with environmental change and increasing aridity at this time (27).

The papionin sample at DMQ (DN 403, 528, 541, 2160, 2162, 2344) is best attributed to *Papio robinsoni*. Diagnostic features include a prominent glabellar region, definitive anteorbital drop, definitive facial fossae, and a dorsally flattened rostrum with rounded maxillary ridges that are elevated superiorly to the nasal bones. They are distinct from *P. hamadryas*, which is found at penecontemporary sites such as Malapa (55), by the aforementioned rostral morphology, rounder maxillary ridges, and by having less excavated facial fossae. However, the specimens collectively share distinct facial fossae that depart from their conspecifics at Swartkrans Member 1. Specifically, the anterior

extent of the maxillary fossae on the DMQ specimens are not as developed as the Swartkrans specimens SK555, SK557, and SK560. The infraorbital region is not as excavated by the maxillary fossae in the DMQ specimens. As the malar root approaches the alveolar bone in the Swartkrans material, it typically curves anteriorly (further defining the maxillary fossae) (56). However, in nearly every DMQ specimen the malar root descends directly on the alveolar bone between the second and third molar. This variation may be explained by temporal variation between DMQ and Swartkrans Member 1, with the latter perhaps dating closer to 1.8 Ma (5). Other sites (e.g., Pit 23 at Bolt's Farm) that have definitively yielded *P. robinsoni* remain undated (57), making DMQ the earliest definitive evidence for this species (56).

The hunting hyaena *Lycyaenops* is first recorded in the early Pliocene of Europe as *L. rhomboideae* (58) and appears by 3.85 - ~3.63 Ma as *L. cf. silberbergi* in the Upper Laetoli Beds at Laetoli (59), and in South Africa at Sterkfontein Members 2 and 4 between 2.61-2.07 Ma (6, 21). During this period *Lycyaenops* cohabits the South African ecosystems with *Chasmaporthetes*, a genus present in South Africa at Langebaanweg by ~5.15 Ma (60-61). DMQ represents the last appearance of *Lycyaenops*, while the South African *C. nitidula* persists into the early Pleistocene (5). This suggests that any environmental conditions favouring the initial dispersal of *Lycyaenops* into South Africa may not have existed after DMQ.

The oldest occurrence of *Dinofelis* (*D. cf. diastemata*) is in South Africa at Langebaanweg ~5.15 Ma (61-62). By the end of the Pliocene *Dinofelis* is represented by *D. aronoki* in eastern Africa, and *D. barlowi* and *D. darti* in South Africa (62). The only *Dinofelis* definitively known to be present in Africa after approximately 1.87 Ma is *D. piveteaui* (62), with *D. barlowi* last known to occur at Malapa and now DMQ at ~2 Ma (9, 32). Moreover, *Dinofelis barlowi* and *Dinofelis* aff. *piveteaui* co-occur at DMQ (29), marking the first time that these taxa have been found together at the same site. Historically, *D. piveteaui* has been interpreted as being directly descended from *D. barlowi* (62-64), but Werdelin & Lewis (62) posited a close relationship between the eastern African *D. aronoki* and the later-occurring *D. piveteaui*. An ancestor-descendent relationship between *D. aronoki* and *D. piveteaui* suggests the expansion of a population of the east African *D. aronoki*, or a population transitional between *D. aronoki* and *D. piveteaui*, at some point

just prior to 2.0 Ma. This population, possibly represented by the DMQ *Dinofelis* aff. *piveteaui* (32) and the Cooper's D *Dinofelis* sp. ~1.36 Ma (21, 65), may have displaced the South African endemic *D. barlowi*.

The dominance of the antilopin *Antidorcas recki* in the DMQ deposits (28.6% of the total bovid assemblage) (32) may reflect movement by an earlier Neogene mammal population from outside South Africa into the evolving palaeocommunities of the region during the earliest Pleistocene. The species is present in the 3.44 Ma Shungura Formation Member B and younger deposits across eastern Africa, ultimately becoming common in terminal Pliocene deposits like the 2.66 Ma Upper Ndolyana Beds at Laetoli (66). In contrast, the first South African *A. recki* specimens are not recovered until the modest sample (NISP: 5, MNI: 3) identified from the Sterkfontein Member 4 assemblage that formed between 2.61-2.07 Ma (with at least some of that fauna post-dating 2.33 Ma due to the occurrence of *Equus*) (6, 26). The hypothesis that South African *A. recki* occurred only ~1 Ma after evolving in eastern Africa, draws support from two potentially interrelated factors. First, *Antidorcas* is notably absent from the large Makapansgat Member 3 assemblage (3.03-2.61 Ma, with faunal deposition close to the end of that period) (6) in which the dominant antilopin and the second-most common bovid is *Gazella vanhoepeni* (NISP: 472, MNI: 55) (23, 66-67). Despite this abundant representation in Member 3, and potential relationship to the Langebaanweg *Gazella* (5.15 Ma) (67) and extant *Gazella* species, *G. vanhoepeni* has not been definitively identified from the extremely sparse record of indeterminate *Gazella* from any subsequent early Pleistocene South African deposit (23, 66, 68-70). Second, shortly after *A. recki* first appears at Sterkfontein Member 4, there is an adaptive radiation of *Antidorcas* in South Africa into at least two additional species (the extinct *A. bondi* and extant *A. marsupialis*). Collectively, species of *Antidorcas* are frequently recovered from early-mid Pleistocene palaeokarstic deposits in the region and form substantial components of the Swartkrans, Kromdraai and Cooper's D faunal assemblages (32, 68-70). The apparent turnover in representation of antilopin genera from *Gazella* to *Antidorcas* coupled with rapid speciation in the latter group is suggestive of exploitation of changing regional palaeoecosystems in which *Antidorcas* species adapted to the progressive expansion of xeric South African Pleistocene palaeohabitats.

## Conclusions and Impact

The geology of DMQ further highlights the inadequacies of the Member system still used at other South African sites (11-14). As DMQ and other sites show (6, 19-20, 70), these different lithologic units can form synchronously, and wherever a siltstone and sandstone unit occurs there is (or was) likely a paired breccia from which it was winnowed. The alternative, allostratigraphic criterion used here disentangles lithostratigraphy from chronology. The context and dating of the DNH 134 *H. aff. erectus* cranium and *P. robustus* fossils to between ~2.04 Ma and 1.95 Ma shows how DMQ is now one of the best dated sites in South Africa, as well as how the integration of geomorphology, stratigraphy, and high-precision age determinations are resolving the South African fossil karstic record and providing critical insights into hominin evolution.

*Australopithecus* is last known to occur ~2 Ma at Malapa (2, 9) and perhaps as late as 2.07 Ma at Sterkfontein (6). *Australopithecus* thus persisted in southern Africa for roughly half a million years after it went extinct in eastern Africa (50). However, the precise timing of the first occurrence of *Homo* and *Paranthropus*, and whether it overlapped with *Australopithecus*, has been hard to resolve because of uncertainty in the depositional ages of fossils from Sterkfontein, Kromdraai and Swartkrans (somewhere between 2.3 and 1.8 Ma). At >1.95 Ma the DNH 152 cranium represents the oldest confirmed representative of *Paranthropus robustus* in South Africa (Fig. 4). The age and association of the Kromdraai B *Paranthropus* fossils are presently unknown, although preliminary palaeomagnetic analysis indicated an age of <1.78 Ma for the hominin-bearing Member 3 deposits (5), and thus younger than DMQ. While the *Paranthropus* fossils from Swartkrans Member 1 Hanging Remnant have been dated to sometime between ~2.3 and 1.8 Ma (5, 21), there is no firm evidence that they are >1.95 Ma (5). A number of studies have suggested the deposits are likely closer in age to the upper flowstone that was dated at 1.8 Ma, consistent with ESR dates (71) and faunal interpretations (72); including the difference noted between *P. robinsoni* at DMQ and Swartkrans (5, 72).

At ~2.04 Ma (minimum age 1.95 Ma) DNH 134 is the most complete and oldest early Pleistocene *Homo* neurocranium (Fig. 2-4) in South Africa. DNH 134 is at least 100-150 ka older than *H. erectus s.l.* specimens from Dmanisi (73) and over 300 ka older than the KNM-ER-3733 cranium from Kenya at ~1.63 Ma (74). The KNM-ER-2598 occipital

fragment from eastern Africa shows affinities to *Homo erectus* (74). However, based on palaeomagnetic data, its location 4m below the KBS tuff (dated to  $1.87 \pm 0.02$  Ma) would place KNM-ER 2598 within the Olduvai Subchron at  $<1.95$  Ma (76-78). As such, DNH 134 also represents the oldest fossil with affinities to *Homo erectus* in the world. Despite this we do not assert that the species necessarily evolved first in southern Africa, especially given major geological biases in hominin finds across Africa. However, the dating of the DNH 134 cranium to  $>1.95$  Ma substantially weakens the hypothesis that *H. erectus sensu lato* evolved outside of Africa (73).

It has been postulated that *A. sediba* is a good candidate for the ancestor of *Homo* (2), although much older fossils attributed to *Homo* exist (8). *A. sediba* can only be ancestral to *Homo* in southern Africa if a population existed prior to DNH 134 for which there is no current evidence (7). An alternative scenario is that the *Homo*-like morphological elements in *A. sediba* (2) may instead represent homoplasy, evolved as local environmental pressures gave rise to convergent morphological features at  $\sim 2.0$  Ma in a terminal population derived from *A. africanus* (79). Regardless, changing environmental conditions across Africa likely placed *Australopithecus* populations under selective pressures that led eventually to the evolution of divergent *Homo* and *Paranthropus* lineages (80-81).

Global climatic transitions have been suggested to occur at 3.0-2.5 Ma (with rapid global cooling and African landscape aridification) and 2.0-1.5 Ma (with the inception of the Walker Circulation in the Pacific governing tropical airflow and rainfall patterns across the continent) (29, 82). Evidence for a major environmental shift in at least parts of southern Africa comes from deep-sea cores off Namibia which indicate stronger climatically induced fluctuations between 2.7 and 2.2 Ma, with semi-arid environments more widespread during glacial periods (27). Aridification and climate variability further increased significantly after 2.2 Ma (27), although humid conditions occurred in the Limpopo basin between 2.0 and 1.75 Ma (28). These climatic shifts are shown in the caves by changes from major phases of massive speleothem and tufa deposition in the Pliocene at the Makapansgat Limeworks and Taung (6), to increasingly sporadic formation throughout the Pleistocene (21). Some researchers have suggested that hominin evolution has been driven by long-term trends towards aridity (83), whereas others suggest the

importance of short periods of extreme climatic change or variability (84). When set within the regional record the DMQ sequence highlights that, while an overall trend towards aridity occurred, the period between 2.3 and 2.0 Ma was a critical period of major ecological change and dispersal that extensively modified South African faunal communities.

The changes seen in hominin taxa are mirrored in other fauna with DMQ representing a transitional faunal community overlapping between older, >2.1 Ma sites like Sterkfontein Member 4 and Makapansgat Limeworks Member 3 (6) and younger sites such as Swartkrans and Gondolin (5). The transitional faunal community recorded at DMQ reflects the larger turnover in mammal communities driven by major ecological changes in southern Africa that saw the extinction of South African endemic species, dispersal of new species, and adaptive radiation in the region. This suggests that South Africa was a centre for the evolution of mammalian lineages as well as a refuge for ancient lineages in the early Pleistocene. The high taxonomic diversity on the South African landscape around 2 Ma is likely a response of indigenous hominins and other fauna to climate and environmental shifts, in tandem with new immigrants that were part of a series of radiations across Africa that set the stage for hominins leaving the continent and inhabiting Asia sometime between 2.12 and 1.95 Ma (85). We interpret the occurrence of *Homo* aff. *erectus* at this time in South Africa, and soon after at Dmanisi (73) as evidence for a major range expansion of this species (covering at least 8000 km) both out of, and within Africa around 2.0 to 1.8 million years ago.

At ~2.04-1.95 Ma DMQ records a critically underrepresented time period in the evolution of South African faunas and palaeoecosystems and shows unequivocally that *Australopithecus sediba* (from Malapa at ~1.98 Ma) (2, 9), *Paranthropus robustus*, and *Homo* aff. *erectus* occurred contemporaneously within the ~250 km<sup>2</sup> of karst landscape NE of Johannesburg, even though it cannot be demonstrated definitively that they were truly sympatric. With the last occurrence of *A. africanus* (Sts 5) potentially as late as ~2.07 Ma at Sterkfontein Member 4 the South African record has a very high diversity of hominins at ~2.1-1.9 Ma (5-6). We suggest that southern Africa served as an ecological refugium for *Australopithecus* until just after 2.0 Ma when either short- or long-term climatic variability would have finally driven *Australopithecus* to extinction. It is unclear whether biological

or behavioural adaptations in *Australopithecus* (80) or competition with *Homo* and/or *Paranthropus* would have also contributed to the demise of the genus, but the dating of DNH 134 and DNH 152 to a period of overlap with *Australopithecus* now make this a possibility.

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**Acknowledgments:** Thanks to the student excavators from the La Trobe University (LTU) Drimolen Palaeoanthropology & Geoarchaeology Field School and University of Victoria Field School, as well as the University of Florence Archaeological Mission to Drimolen. In particular we thank LTU PhD student Richard Curtis who first found parts of the DNH 134 cranium. We thank D. and J. Smith and Khethi Nkosi, the landowners at Drimolen at various stages, and who granted our permission to work at the site in conjunction with the South African Heritage Resources Agency (SAHRA). The DNH 152 cranium is named after the landowner (Khethi) as he excavated with us on the field school and discovered the first tooth. We want to say a special thanks to our co-author Simon Mokobane for his many years of work at the site. Simon recently passed away due to cancer and he will be remembered fondly by many generations of people who have worked at Drimolen. In recognition of this we have named the DNH 134 cranium after him (Simon). This work would not have been possible without him. We thank Dario Bilardello, Mike Jackson and Joshua Feinberg for their help and advice while measuring at the Institute of Rock Magnetism. Thanks to Bernhard Zipfel for facilitating access to the hominin collections at the University of the Witwatersrand. **Funding:** The bulk of this research was funded by Australian Research Council Future Fellowship Grant FT120100399 to AIRH and ARC Discovery Grant DP170100056 to AIRH, JWA, DSS, and RJB. The U-Pb analysis was funded by ARC DECRA DE120102504 to RP. The US-ESR dating was supported by ARC DP140100919 to RJB. Work at the site by the Italian Archaeological Mission was supported by a series of grants by the Italian Ministry of Foreign Affairs to JMC. CM thanks the National Research Foundation (African Origins Platform) for grants that supported the excavation and research at Drimolen. This work was also supported by a La Trobe University Postgraduate Research Scholarship and La Trobe University Internal Research grant to ABL, AM, BA, TE, TM and RS, and a Society of Antiquaries London research grant to JMM. Components of the palaeomagnetic work were conducted during a Visiting Research Fellowship to TM at the Institute for Rock Magnetism, University of Minnesota, supported through the National Science Foundation, USA. **Author contributions:** AIRH led the geology and dating program and undertook the palaeomagnetic analysis with TM. JMM, ABL, DS, SN, GS and CM undertook the analysis of the hominin remains. JWA, JM, DSR, CM and SB undertook faunal analysis and

biochronology. GB and AIRH undertook stratigraphic analysis and interpretation. GB, AM, TE and TD undertook micromorphological analysis. RJB undertook the US-ESR dating. BA undertook spatial and GIS analysis. JH, JW and RP undertook uranium-lead dating and associated contextual analysis. MC and RCS undertook studies of the archaeological material. ABL, SB, JMM and AIRH excavated the hominin crania. BJA and CW undertook video and imaging of the site. AIRH, SB, CM, and JMC directed excavations at different times that recovered the hominin material. CM initiated the dating program at Drimolen. All authors contributed to the writing of the paper. Research supporting this publication was undertaken by JM, ABL, AM, while completing a PhD at La Trobe University **Competing interests:** Authors declare no competing interests. **Data and materials availability:** All data are available in the main text or the supplementary materials. The fossils are curated at the University of the Witwatersrand hominin vault. Access for research is granted through the hominin access committee ([Bernhard.Zipfel@wits.ac.za](mailto:Bernhard.Zipfel@wits.ac.za)), including requests for the 3D scan data.

**Fig. 1. The Location of the Drimolen Palaeocave Complex.** (A): The location of Drimolen in comparison to other *Homo erectus* sites worldwide and their approximate maximal age. (B) The location of Drimolen within South Africa in relation to other early hominin fossil sites and the Gauteng Malmani sites. (C) The location of Drimolen within the Gauteng Malmani in relation to other hominin sites. The Blauubank Stream Valley is represented by sites running from Bolt's Farm (BF) to Kromdraai (KR) (CP, Coopers D; STK, Sterkfontein; SWT, Swartkrans; RS, Rising Star; GV, Gladysvale; ML, Malapa; HG, Haasgat; GD, Gondolin). Colours indicate the predominant genus or species represented (C) An aerial view of the Drimolen site and the relationship of the hominin bearing DMQ (2.04-1.95 Ma) and non-hominin bearing DMK (~2.61 Ma).

**Fig. 2. The DNH 134 *Homo aff. erectus* neurocranium.** (A) Superior view, anterior to the left. (B) Posterior view. (C) Right lateral view, anterior to right. (D) Left lateral view, anterior to the left. Scale bar = 10 mm.

**Fig. 3. Endocranial volume estimation of DNH 134.** (A) Endocranial landmark set used for ECV estimation. Each vertex of the surface is used as a landmark or semilandmark. Anatomical landmarks are shown as spheres, curve semilandmarks are connected as black lines. Measured versus predicted ECV [(B) regression-based, (C) pooled TPS-based, and (D) species-specific TPS-based estimates] for human (blue), *H. erectus* (red), gorillas (grey), orangutans (orange), chimpanzees (green).

**Fig 4. Comparisons of the DNH 134 cranium with the Mojokerto juvenile *H. erectus* cranium.** DNH 134 (red) superimposed on the Mojokerto cranium (grey) after scaling both specimens to the same bregma – inion length. (A) Left lateral view, anterior to left. (B) Superior view, anterior to left. (C) Posterior view.

**Fig. 5. The DNH 152 *Paranthropus robustus* cranium.** (A) Superior, (B) Posterior, (C) Right Lateral, (D) Left Lateral, (E) Right Temporal, (F) Right Orbit, (G) Left M1 Occlusal, (H) Left M2 Occlusal, (I) Left M1 Buccal, (J) Left M2 Buccal, (K) Right M1 and M2 Buccal, (L) Right M1 and M2 Occlusal. Scale bar = 10mm

**Fig. 6 3D laser scan of Drimolen Main Quarry.** The location of the hominin fossils DNH 7, DNH 134 and DNH 152 are shown relative to the main stratigraphic sections (Warthog Cave, Italian Job, Jangi, Walls of Jericho) and features described in the text (WC, Warthog Cave; CEA, Central Excavation Area; WW, dolomite Western Wall of the palaeocavern).

**Figure 7. Warthog Cave Section** showing palaeomagnetic and uranium-lead (U-Pb) and sample locations: 1: dolomite bedrock; 2: flowstone (BFS); 3: clast-supported breccia (CSB); 4: sandstone/siltstone (WGSS); 5: recent decalcified sediment, mine dumps and displaced blocks; 6: U-Pb sample locations: BFS (DN39A;  $2.673 \pm 0.103$  Ma) and WCFS (DN09  $1.789 \pm 0.104$  Ma). Palaeomagnetic samples indicated by black (normal polarity) and white (reversed polarity) circles

**Figure 8. Jangi Buttress Section** depicting DNH 134 find locus: JB - Jangi Buttress; EB - Eurydice Block; 1. CSB, with skeleton-supported structure, large chert and decalcified dolomite (grey dusty patches) blocks and preserved, microfauna-rich reddish layers; 2. MSB with unsorted chert and subordinate dolomite clasts (the lower half is decalcified); 3. mining rubble filling void cut underneath Jangi Buttress; dotted rectangle depicts DNH 134 *Homo* cranium dispersion area; DMQ-2 denotes the US-ESR date  $2.041 \pm 0.240$  Ma (50cm scale).

**Figure 9. Italian Job Pinnacle.** Photograph (left) and Section (right) showing stratigraphy and geochronometric data: 1. dolomite bedrock; 2. MSB; 3. flowstone (WOJFS); 4. GSS; 5. decalcified sediment; 6. DMQ-3 denotes the US-ESR date  $1.962 \pm 0.107$  Ma; red circle - DNH 152 *Paranthropus* skull location.

**Figure 10. Walls of Jericho Pinnacle.** Photograph (left) and Section (right) showing stratigraphy, palaeomagnetic polarity and geochronometric data: 1. dolomite bedrock; 2. wad and other dolomite weathering products; 3. basal flowstone (BFS); 4. flowstone; 5. GSS: sandstone/siltstone; 6. GSS: coarse sandstone /fine gravel; 7. decalcified sediment; 8. Walls of Jericho Flowstone (WOJFS) dated by U-Pb  $1.962 \pm 0.107$  Ma. Palaeomagnetic data: black circles (normal polarity) grey circles (intermediate polarity); white circles (reverse polarity).

**Figure 11. Micromorphology of selected aspects of the DMQ infill.** (A) HR scan of thin section of CSB (main talus cone) sample MM26, from WW, showing loose microstructure (LMs), chert (Ch) and dolomite (Ds) clasts, reworked soil/sediment aggregates (SA) sometimes showing crust-like features probably reworked from sandstone/siltstone sediments. (B) CSB sample MM04 from WW, with clay aggregate stained by amorphous Fe- and subordinately Mn-oxides; the lower third of the panel is under XPL, showing poorly developed stipple-speckled b-fabric (white arrows) in unstained areas. (C) CSB sample MM01 from WC, with bone fragment (Bf), monocoth-like wood fragment (white arrow), wide pores originated by biological activity (BV); the lower third of the panel is under XPL, showing anhedral sparite infills within pores. (D) Sample MM32 from WC; void with “dusty” calcite coating (CC) probably deriving from recrystallisation of aragonite, and subsequent sparitic anhedral calcite infilling (CI). (E) HR scan of thin section of MSB (intermediate facies) sample MM18L from IJ; sequence of fining-upwards sequences with cm-size clasts at the base of each sequence; the bracket highlights the sequence in panel F. (F) Mosaic of microphotographs under PPL (left) and XPL (right), showing a f.u. sequence with thin clay crusts interbedded within the fine part. (G) HR scan of thin section of GSS (distal facies) sample MM12 from WOJ (above WOJFS), with fine grainsize fining-upwards sequences (black brackets), interbedded with a thin flowstone crust (FS); the red and yellow squares indicate respectively the areas in panels H and I. (H) very fine silt and clay crusts (Cr) topping f.u. sequences. (I) Columnar calcite flowstone (Fs) with multiple short growth hiatus marked with detrital caps on crystal tips. the flowstone is overlain by a f.u. sequence terminated by a clay crust (white arrow); the right side of the panel is under XPL.

**Figure 12. Micromorphology of the Wall of Jericho Flowstone.** (A) aspect and stratigraphy of the ~1.95 Ma WOJFS, as in the northern side of the Walls of Jericho pinnacle, showing that WOJFS formed during a stop in clastic deposition and is not intrusive into GSS; FBU1 and FBU2: Flowstone-Bounded Units 1 and 2, as in Figure 4.; yellow rectangle: micromorphology samples. Scale = 5cm, (B) short range picture of the micromorphology sample detachment niche inside red rectangle of panel A. BFS: basal flowstone, (C) scan of thin section showing the distinct nature of sedimentation below and above WOJFS, with several fine speleothem crusts alternating with silt before the formation of the main flowstone. Coloured rectangles indicate spots described in the following panels; blue: panel D, green: E, amber: F, black: G; red: H, (D) top of WOJFS and upper contact (white arrow) with the overlying USS. Black arrow: subhedral sparitic calcite with rhombohedron faces indicating upwards crystal growth direction. The layer of anhedral calcite crystals between the arrows is recrystallised and the top surface has undergone dissolution due to contact with siltstone. Plane Polarised Light (PPL), (E) same as in D, showing a wide gulf-like dissolution feature (DF) into the upper surface of WOJFS layer, due to contact with subsequently deposited silt. White arrows indicate remains of the upper layer (PPL), (F) high magnification of calcite within WOJFS, showing the remnant needle-like aragonite (black arrows) that was critical to the successful dating of the sample by U-Pb (PPL), (G) Thin flowstone crust (FS) underlying the main WOJFS, showing preserved upwards-growth pattern. Lens-like voids were subsequently infilled by anhedral calcite (CIV), (H) Precipitation of anhedral calcite spar (CIV) within channel voids.

**Fig. 13: Palaeomagnetic data from DMQ.** (Upper) Representative palaeomagnetic data plots for DMQ (vector, stereographic, and demagnetisation spectra). Open symbols on stereographic plots = negative inclination and closed symbols = positive inclination. (A) DN09 Normal Polarity >2.61 Ma Basal Flowstone, (B) DN27 Normal Polarity Siltstone from the top of the Walls of Jericho, (C) DNFS3 Intermediate Polarity ~1.95 Ma Walls of Jericho Flowstone, (D) DN29 Reversed Polarity Siltstone from base of the Walls of Jericho, (E) DN01 Reversed polarity siltstone from the Warthog Cave Section, (F) DN10 Rejected Basal Flowstone Sample with normal trend but with MAD >15. (Lower) Mineral magnetic results for the DMQ. (G) Unmixed coercivity contributions to a representative backfield curve with labelled remanence coercivities for each component (comp.). (H) FORC diagram with smoothing parameters listed (e.g., Sc0) and a hysteresis loop insert. (I) thermomagnetic curve with curie temperature estimate.

**Fig. 14. Composite stratigraphy for DMQ.** The Stratigraphic Sections (WC – Warthog Cave, JB – Jangi Buttress, IJ – Italian Job, WOJ – Walls of Jericho; m = metres below or above datum) and Dates (Ma) for DMQ compared against the Geomagnetic Polarity timescale and other early hominins in South Africa as well as global *Homo erectus* sites.

SAMPLE	DMQ-2	DMQ-3
<b>ENAMEL</b>		
Dose (Gy) <sup>a</sup>	1814±59	2414±76
U (ppm) <sup>b</sup>	1.43±0.15	2.07±0.18
<sup>234</sup> U/ <sup>238</sup> U <sup>b</sup>	1.2841±0.0587	1.0886±0.0369
<sup>230</sup> Th/ <sup>234</sup> U <sup>b</sup>	0.9220±0.0256	0.9731±0.0155
Thickness (m)	1354±189	1551±320
Water (%)	3±1	3±1
<b>DENTINE</b>		
U (ppm) <sup>b</sup>	17.30±1.11	19.31±1.05
<sup>234</sup> U/ <sup>238</sup> U <sup>b</sup>	1.4110±0.0237	1.4183±0.0109
<sup>230</sup> Th/ <sup>234</sup> U <sup>b</sup>	0.9558±0.0201	0.8822±0.0278
Water (%)	5±3	5±3
<b>SEDIMENT</b>		
U (ppm)	1.9±0.5	1.9±0.5
Th (ppm)	3.02±0.2	3.02±0.2
K (%)	0.29±0.05	0.29±0.05
Water (%)	15±10	15±10
<b>EXTERNAL DOSE RATE SEDIMENT</b>		
Beta dose (μGy a <sup>-1</sup> )	44±9	38±8
Gamma Dose (μGy a <sup>-1</sup> )	323±39	323±39
Cosmic (μGy a <sup>-1</sup> )	97±50	97±50
<b>COMBINE US-ESR AGE</b>		
Internal dose rate (μGy a <sup>-1</sup> ) <sup>c</sup>	291±80	701±76
Beta dose dentine (μGy a <sup>-1</sup> ) <sup>c</sup>	134±37	69±8
P enamel <sup>c</sup>	0.64±0.13	-0.45±0.01
P dentine <sup>c</sup>	0.03±0.02	1.46±0.15
Total dose rate (μGy a <sup>-1</sup> ) <sup>c</sup>	889±109	1228±100
<b>AGE (ka)<sup>c</sup></b>	<b>2041±240</b>	<b>1965±147</b>

**Table 1. US-ESR dating data for DMQ-2 and DMQ-3 fragments**

<sup>a</sup> Dose equivalent  $D_e$  obtained using *McDoseE 2.0*, with *SSE* (from Joannes-Boyau et al<sup>101</sup>).

<sup>b</sup> Uranium concentration values were obtained by LA-MC-ICPMS and LA-ICPMS on both teeth and both dental tissues.

<sup>c</sup> The age was calculated using Shao et al<sup>50</sup>, with the dose rate conversion factors of Guérin et al<sup>110</sup>, the enamel and dentine density of 2.95 and 2.85 respectively from Grun<sup>110</sup>.



**Table 2. Palaeomagnetic data, associated US-ESR and U-Pb ages and age ranges for the various deposits at DMQ. (FS = flowstone, SS = sandstone and siltstone)**

Sample	Location	Depth	Type	Dec	Inc	K	Plat	Polarity	U-Pb/ESR (Ma)	Combined Date (Ma)	Fossil
DN09	WC	-3.90	FS	16.5	-63.7	141.4	66.5	N	1.79 ± 0.10	1.89-1.78	
DN27	WOJ	0.04	SS	11.7	-23.2	40.3	72.3	N		1.95-1.78	
DN24	WOJ	-0.45	SS	16.9	-21.0	149.3	68.0	N		1.95-1.78	
DN26	WOJ	-0.59	SS	42.7	-12.0	64.5	44.7	I		~1.95	
DNFS3	WOJ	-0.78	FS	242.1	-53.8	62.2	-5.8	I	1.96 ± 0.11	~1.95	
DN19	WOJ	-1.02	SS	222.3	28.6	174.7	-48.4	I	1.97 ± 0.15	~1.95	
DN21	WOJ	-1.55	SS	258.2	29.4	337.6	-17.0	I		~1.95	
DN29	WOJ	-1.71	SS	153.8	38.3	41.4	-65.7	R		2.28-1.95	
DN08	WOJ	-1.90	SS	156.9	29.6	56.0	-61.6	R		2.28-1.95	
DN01	WC	-3.33	SS	167.4	17.1	139.0	-69.0	R		2.28-1.95	DNH 152
DN03	WC	-4.36	SS	188.4	26.7	180.7	-75.8	R	2.04 ± 0.24	2.28-1.95	DNH 134
DN39	WC	-6.40- 80	FS	26.5	-29.1	156.9	63.1	N	2.67 ± 0.10	2.77-2.61	

## Materials and Methods

### Reconstruction of endocranial volume

Based on surface scans of the original fossil acquired through use of the Artec Spider scanner following methodology outlined in Adams et al. (86), we extracted the preserved endocranial surface of DNH 134 and generated a partial virtual endocast by smoothly filling in missing regions. We used a symmetric human template of an endocranial landmark configuration (935 landmarks and sliding semilandmarks ; Fig. 3) (see also 87-88), and defined which landmarks and sliding semilandmarks are preserved and which are missing. Semilandmarks were slid to the symmetric template configuration to gain point-to-point correspondence between individuals using the bending energy of the thin-plate-spline algorithm as a minimization criterion. Endocranial (semi)landmarks were also captured for an extant reference sample including 90 humans, 27 chimps, 39 gorillas and 44 orangutans, as well as for four African and four Asian *H. erectus* individuals (KNM-ER-3733, KNM-ER-3883, KNM-WT-15000, OH9, Sambungmacan 3, Sangiran 2, Ngandong 14, Nwagi) (89).

On the basis of the extant reference sample, a multiple linear regression model was established to estimate endocranial volume from the endocranial form of the preserved regions of DNH 134 as captured from the landmarks (e.g., 89). Pretending that the same regions as in DNH 134 are missing in each of the reference individuals as well as the listed *H. erectus* individuals and using the same methodology to estimate their endocranial volume allows comparison of the estimated and actual endocranial volume and therefore

an interpretation of how the choice of the reference sample influences the estimates (e.g., 89-90). This analysis (Fig. 3) shows that predicted and actual endocranial volumes are highly correlated without a taxon-dependent bias towards over- or underestimation so that DNH 134's endocranial volume can be estimated reliably.

Additionally, the missing portions of DNH 134's endocranial surface were reconstructed based on thin-plate-spline warping of the extant reference sample and the reconstructions' endocranial volumes were measured. The average or most common value and the range of estimates based on different reference individuals can be interpreted as the most probable value and estimation uncertainty, respectively (90-91). Assuming that the same regions as in DNH 134 are missing in each of the reference individuals as well as the listed *H. erectus* individuals, we estimated their endocranial volumes as well and compared them to their measured cranial capacities (Fig. 3). If the pooled reference sample including humans and apes is used, the endocranial volume of smaller individuals (apes) is predicted too small. A good estimate is possible if the "correct" reference sample (i.e. the same species) was used (Fig. 3). Using humans and apes in a pooled reference sample to reconstruct DNH 134 might therefore lead to (slight) over- or underestimation of its endocranial volume. However, our results show that endocranial volume estimates based on thin-plate spline reconstructions (484-593 cc) are consistent with regression-based estimates (514-564 cc).

As a two-to-three-year-old individual, DNH 134 has not reached its adult brain size when it died, although it was likely approaching the adult population range of variation. We computed Gompertz growth curves for ontogenetic data of human and chimpanzee data from the literature for comparison (87, 92-97). A two-year-old individual with a

cranial capacity of 538 cc would grow into an adult of 661 cc and 577 cc according to the human and the chimpanzee growth curve, respectively, while a three-year-old individual would grow into an adult of 588 cc and 551, respectively. Despite its juvenile status, we therefore suggest that DNH 134 documents small brain size in this population.

### **Stratigraphy and Micromorphology**

The characteristics of the sediments were observed and described on natural profiles (dissolution shapes), old quarry cuts, and recent excavation profiles and surfaces. Descriptions were carried out following Catt (98). Emphasis was given to texture, boundaries, sedimentary structures and, more generally, to the architecture of the sedimentary bodies, to reconstruct the stratigraphy of the cave infill. Litho- and allostratigraphic units (99) were used in assessing the stratigraphic sequence and then lumped into chronostratigraphic Flowstone Bound Units (22). Texture data were inferred from field observations and the measurement of clasts in exposed profiles, on freshly broken sections of hand samples, and from micromorphological samples in thin section under polarising microscope. Micromorphological observations were carried out on undisturbed sediment samples. Oriented blocklets were detached from natural or excavation profiles; their coordinates were recorded by a total station, and by far-field and macrophotography. The samples were air-dried in laboratory at 35°C in ventilated oven, and impregnated, including the cemented samples, by epoxy resin under medium vacuum and let polymerise. The resulting samples were cut by diamond saw, polished, glued on 90x60 or 60x45 mm microscope slides and ground to 30 µm on abrasive disks, using petroleum for lubrication. The resulting slides were protected by standard cover slides and

labelled. Observations were carried out following the standard formalised by Bullock et al (100) and Stoops (44). Thirty-eight monoliths were collected from locations selected to represent the most relevant aspects of the various lithologic units, as well as local peculiarities.

### **Coupled Uranium series and Electron Spin Resonance (US-ESR)**

For each tooth a small fragment of dentine and enamel was removed using a hand-held saw, with a 300 µm thick diamond blade. The small fragment was then cut in half exposing a flat surface of dentine and enamel for U-series analyses. Internal dose rate was calculated using U-series values obtained via LA-MC-ICPMS on a Thermo Neptune plus coupled to ESI NW193 and ICPMS quadrupole Agilent 7700 for concentration correction of  $^{238}\text{U}$ . The values measured in the enamel and the dentine, are consistent across the measured area. An average value was calculated for each dental tissue and used for the US-ESR model. Baseline and drift were corrected using a NIST 612 glass disc, while a fossil hippopotamus tooth of known U-series concentration was used to correct  $^{234}\text{U}/^{238}\text{U}$  and  $^{230}\text{U}/^{238}\text{U}$  ratios and assess the accuracy of measurements. Concentration obtained by LA-MC-ICPMS were compared to quadrupole ICPMS analyses on the same dentine and enamel fragments (Table S2). To account for tailing effects, measurements were carried out at half-masses of 229.5 and 230.5 for  $^{230}\text{Th}$  and 233.5 and 234.5 for  $^{234}\text{U}$ . The other half was used for ESR measurements, by separating the different dental tissues. Both DMQ-2 and DMQ-3 enamel fragments removed from each sample and used for the ESR measurements were directly in contact with the dentine on one side and directly in contact with the sediment on the other side (no cement). The outer surface of the enamel (in contact

with the surrounding sediment) and the dentine directly attached to it were removed using a diamond blade rotary tool. Simultaneously, 100  $\mu\text{m}$  on each side was removed to avoid alpha particle contribution.

ESR dating for Drimolen fossil teeth was performed on a Freiberg MS5000 X-band spectrometer at 1G modulation amplitude, 2mW power, 100G sweep, and 100KHz modulation frequency, coupled to a Freiberg X-ray irradiation chamber, which contains a Varian VF50 x-ray gun at a voltage of 40KV and 0.5mA current. Each tooth fragment was mounted onto a teflon sample holder, allowing the fragment to be exposed directly to the x-ray source with no shielding (except for a 200 $\mu\text{m}$  aluminium cover). To estimate the ESR equivalent dose ( $D_E$ ), each fragment was irradiated nine times, following exponentially increasing irradiation times (i.e. 90s, 380s, 1080s, 1800s, 3600s, 7200s, 14700s, 28800s and 63300s, with an average dose rate for DMQ-2 and DMQ-3 of 0.22Gy/s 0.25Gy/s respectively). The x-ray emission received by the bovid teeth was calibrated using added known gamma irradiation dose performed at the Australia's Nuclear Science and Technology Organisation (ANSTO). During each irradiation step, the output of the x-ray gun was recorded, to allow an accurate determination of the dose received by the sample at each irradiation steps. Fitting procedures were carried out with the *McDoseE 2.0* software that uses a Bayesian framework approach, where the solution is a full probability distribution on the dose equivalent (101). The dose response curves (DRCs) were obtained by averaging the peak-to-peak  $T_1$ - $B_2$  ESR intensities recorded for each irradiation dose over 180 degree (10° step) measurements and merged into a single spectrum (47). Isotropic and baseline corrections were applied uniformly across the measured spectra (48). The final  $D_E$  values were obtained by fitting a single saturating exponential (SSE) through the ESR

intensities and by selecting the appropriate maximum irradiation dose ( $D_{\max}$ ) in order to avoid dose estimation inaccuracy (102).

The external dose rate (Table S3) was calculated using the U, Th and K content of sediment collected from the site as well as measured directly using a portable gamma-spectrometer Inspector 1000. Measurements were made in places where breccia was already fractured. Some results where the probe could not be fully surrounded by 30cm of sediment were modelled using the sedimentary geometric configuration around the probe (*Modelled breccia*), but represent poorly constraint measurements and therefore removed from calculation. The Drimolen cosmic dose rate was estimated, considering the site variation over the burial time as detailed in this study. With an estimated cover of about 30m of dolomite (using  $2.85 \pm 0.03 \text{ g.cm}^{-3}$  average density value) in the original context and a gradual denudation rate of 10 meters per million years. The large error takes into consideration sudden collapse of parts of the dolomite instead of a gradual denudation rate (Table S3).

### **Palaeomagnetism**

A series of 11 independently oriented samples (Table 2) were taken from the Drimolen palaeocave deposits targeting clastic siltstone sequences and flowstones. These were taken as block samples and were oriented *in situ* using a Suunto magnetic compass and clinometer. Subsequent corrections were made for the local dip of the stratigraphy and the declination of the local field according to the International Geomagnetic Reference Field (IGRF), accessed through the British Geological Survey ([http://www.geomag.bgs.ac.uk/data\\_service/models\\_compass/igrf\\_form.shtml](http://www.geomag.bgs.ac.uk/data_service/models_compass/igrf_form.shtml)). Block

samples were drilled and cut into standard 20 x 25 mm palaeomagnetic subsamples using water-cooled preparation equipment. Palaeomagnetic experiments were undertaken at TAAL, with additional mineral magnetic tests performed at the University of Liverpool Geomagnetism Laboratory (ULGL; UK) and Institute for Rock Magnetism (IFRM), University of Minnesota (US). Mass specific magnetic susceptibility measurements at low ( $\chi_{LF}$ ) and high ( $\chi_{HF}$ ) frequency were undertaken using a Bartington MS3 system for frequency-dependant, room-temperature analysis. Isothermal Remanent Magnetization (IRM) acquisition curves and backfields, hysteresis loops and thermomagnetic curves ( $M/T$ ) were run on a Magnetic Measurements Variable Field Translation Balance (VFTB) at ULGL, with additional IRMs imparted using a MMPM10 Pulse Magnetiser at TAAL. Curie temperatures ( $T_c$ ) were calculated from  $M/T$  heating curves using the Moskowitz et al. protocol (103) smoothed with a 3-point running average in RockMagAnalyzer 1.1 (104). First-order reversal curves (FORCs) were measured on a Princetown Micromag Vibrating Sample Magnetometer at IFRM and processed using FORCinel 3.0 (104) and VARIFORC smoothing (106). Hysteresis backfield curves were de-convoluted using MAX UnMix (107). Palaeomagnetic samples primarily underwent a 16-point thermal demagnetisation ( $TH_D$ ) using a shielded Magnetic Measurements MMTD80a Thermal Demagnetiser in a zero-field cage. Alternating field demagnetisation ( $AF_D$ ) was also undertaken for comparative purposes using a Molespin alternating field demagnetiser and an Advanced Geoscience Instruments Company (AGICO) LDA5 Alternating Field Demagnetizer. An additional hybrid demagnetisation strategy was employed incorporating low field  $AF_D$  steps (e.g., to 8–12 mT) prior to standard  $TH_D$  which has shown to be useful in removing viscous overprints in palaeokarst deposits (9, 18-19). All remanence



measurements were made using an AGICO JR6 spinner magnetometer at TAAL. Subsample characteristic remanent magnetisation (ChRM) directions were isolated using principle component analysis (*108*) and were accepted with median angle of deflections (MAD) of  $<15^\circ$ . Final directions for each block sample were calculated from between 3 and 7 subsample ChRMs using Fisher (*109*) statistics. Polarity directions were assigned based on virtual geomagnetic pole latitudes of +90 to +60 (normal), +60 to -60 (intermediate), and -60 to -90 (reversed).

### **Supplementary Material**

Tables S1-S3.

Movie S1.